

Recent Developments at ECMWF

Nils Wedi

Thanks to Martin Miller, Agathe Untch, Peter Bechtold,
Dick Dee and Mats Hamrud for their contributions to these slides.

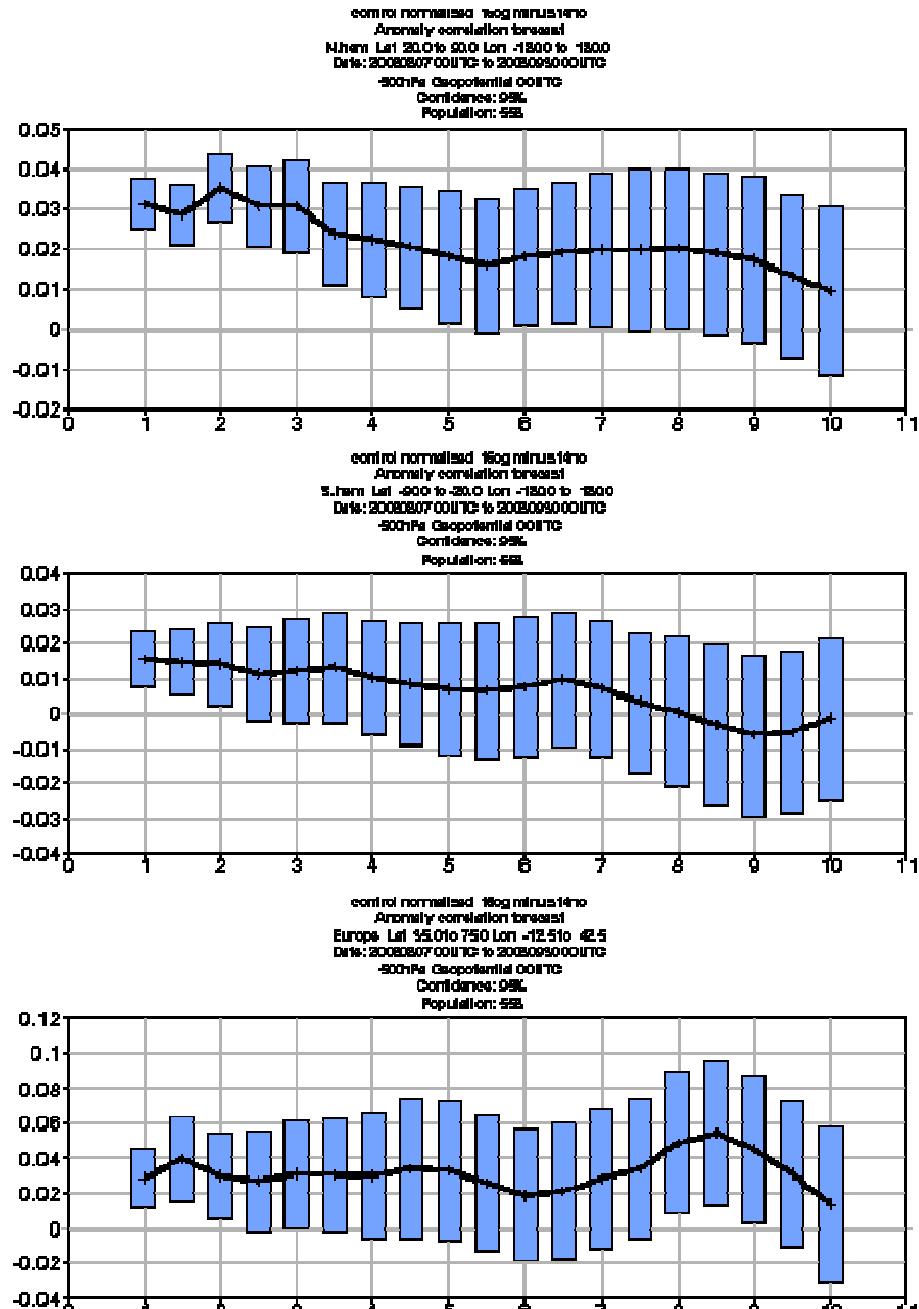
Overview

- ◆ The current operational cycle CY35R3
- ◆ Improvements in the Stratosphere
- ◆ 1989-2009 ERA-Interim
- ◆ The next high resolution implementation T_L1279
- ◆ Testing of the non-hydrostatic IFS

The current operational cycle 35r3

- ◆ Non-orographic gravity wave scheme
- ◆ New trace gas climatology
- ◆ *Revision of the snow scheme*
- ◆ *New structure of the surface analysis*
- ◆ *Improved assimilation of land-surface sensitive channels*
- ◆ Assimilation of total column ozone data from METOP; water vapour data from MERIS; cloud-affected radiances for infrared instruments
- ◆ Various changes to 4D-VAR: new humidity formulation in 4D-Var, improved quality control (for conventional obs), weak-constraint 4D-Var (stratospheric model error)

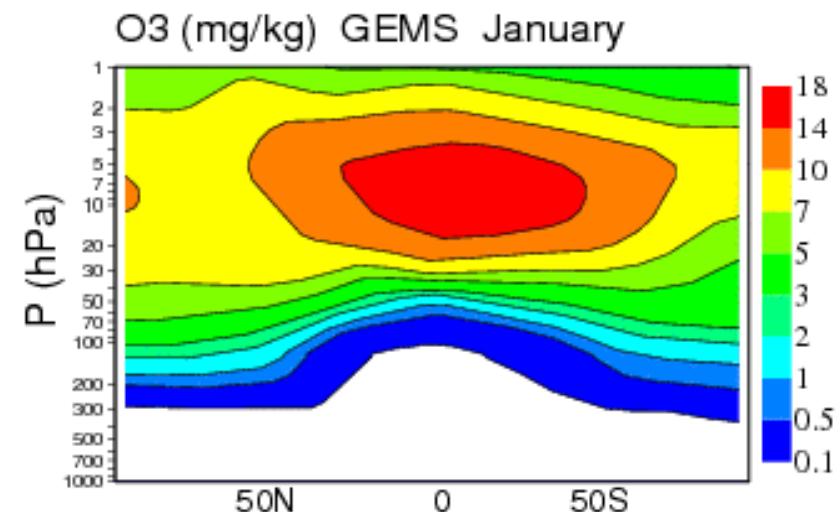
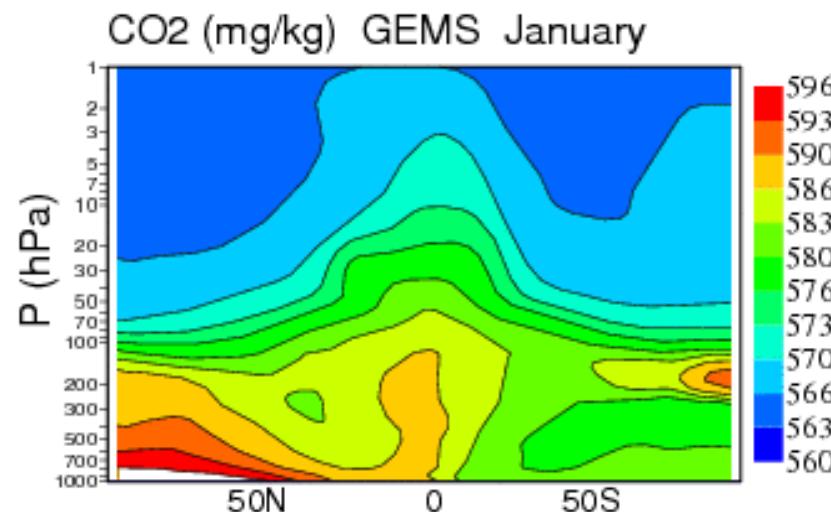
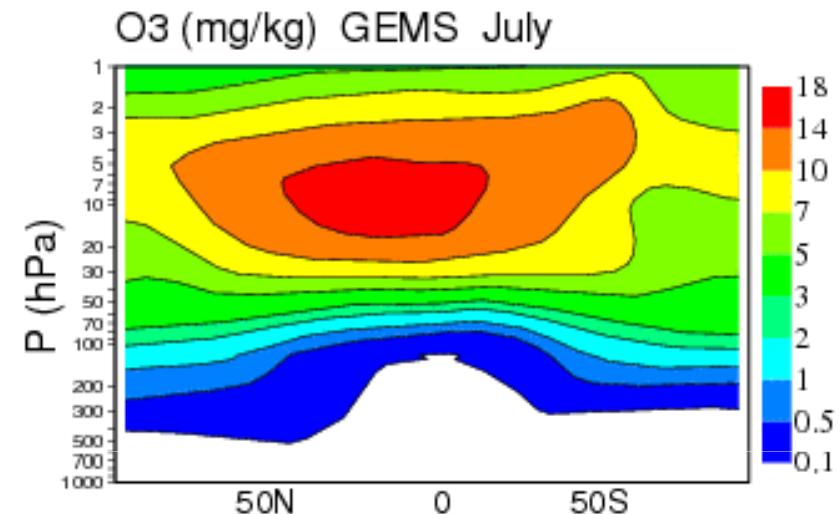
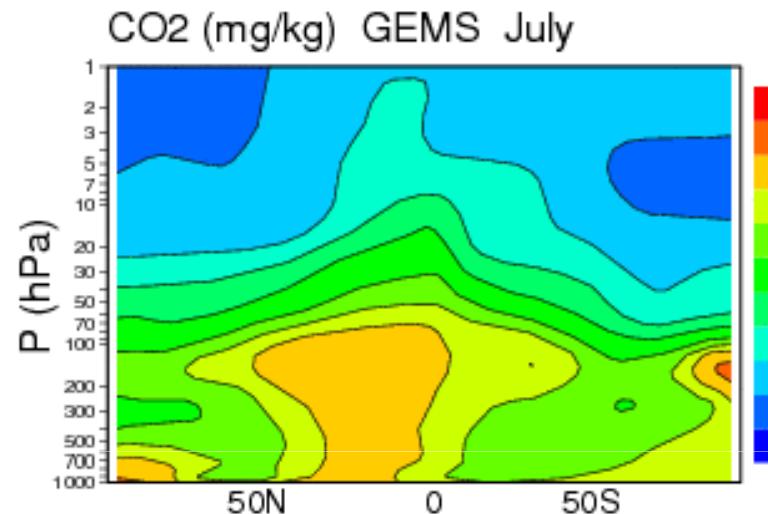
- ◆ CY35R3 anomaly correlation between forecast and operational analysis deviations from climate with statistical significance based on 95% confidence interval.
- ◆ Extensive testing: 558 days of assimilation and forecast!



Improvements in the Stratosphere

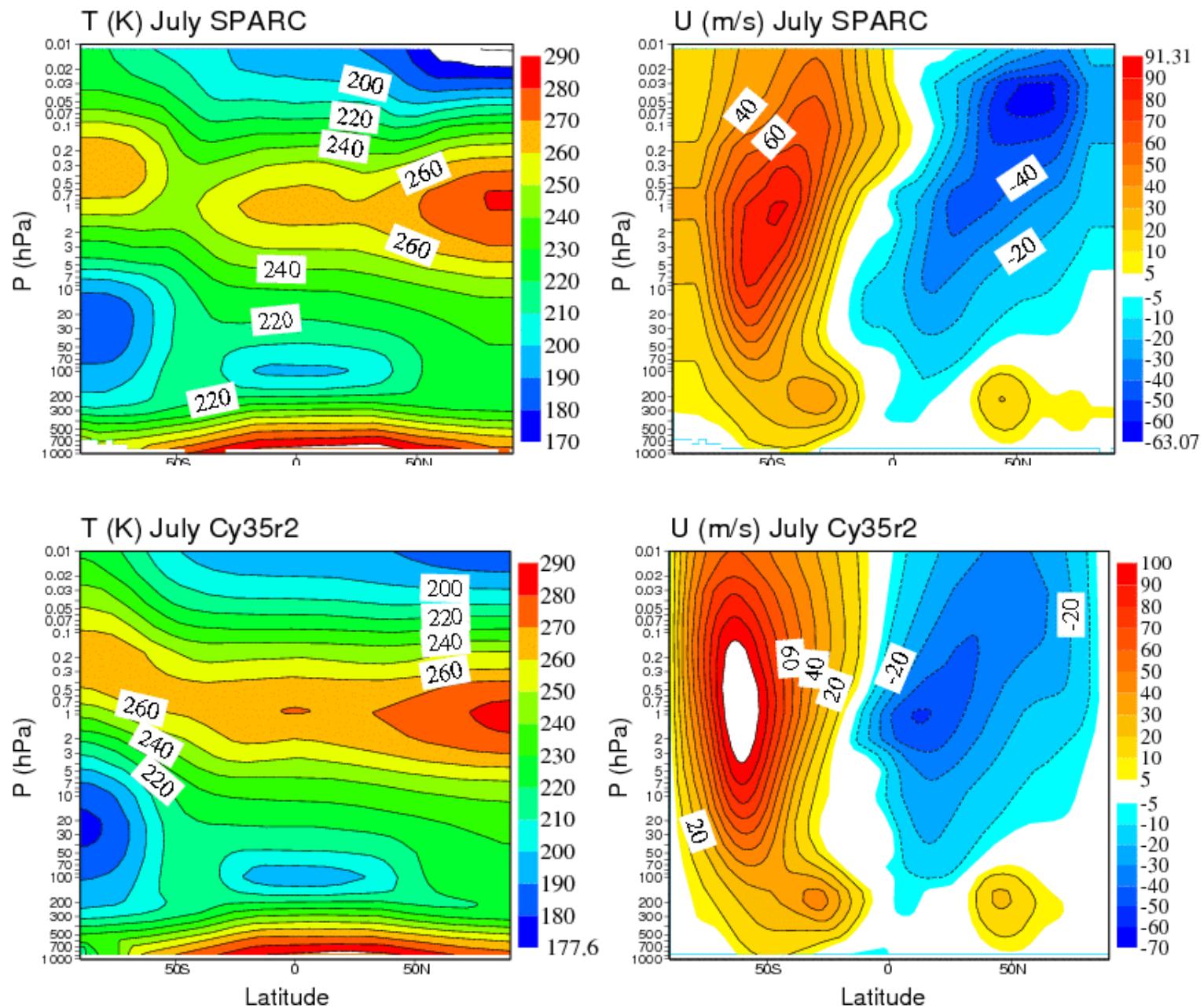
- ◆ New trace gas climatology
- ◆ Non-orographic gravity wave scheme (*Warner and McIntyre, 1996; Scinocca, 2003, ECMWF Newsletter 2009*)
 - ◆ Define launch height (450hPa, *Ern et al., 2006*)
 - ◆ *Define amplitude and shape of the launch spectrum of waves with various vertical wavenumbers, propagating directions + horizontal phase speeds*
 - ◆ *Dissipation mechanisms: critical level filtering by the prevailing (model) winds and nonlinear dissipation due to wave steepening and breaking*

CO₂ and O₃ zonal mean concentrations from GEMS as used in Cy35r3



SPARC

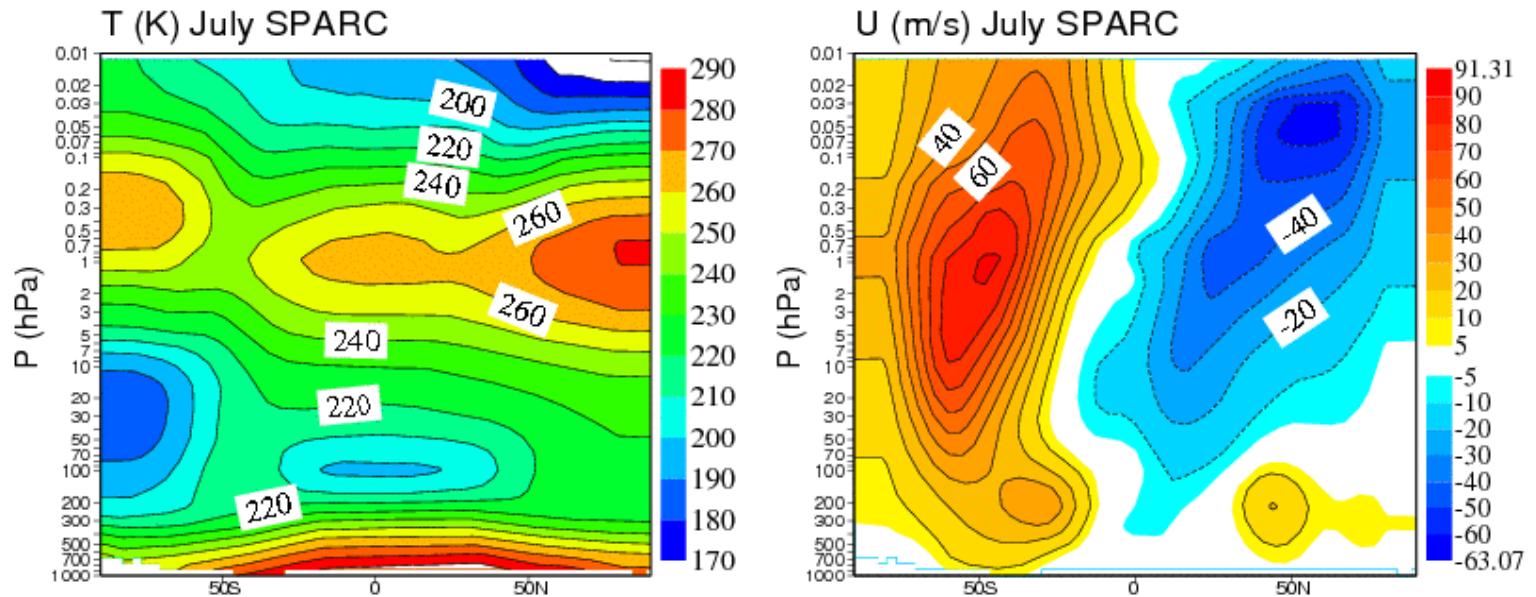
July climatology



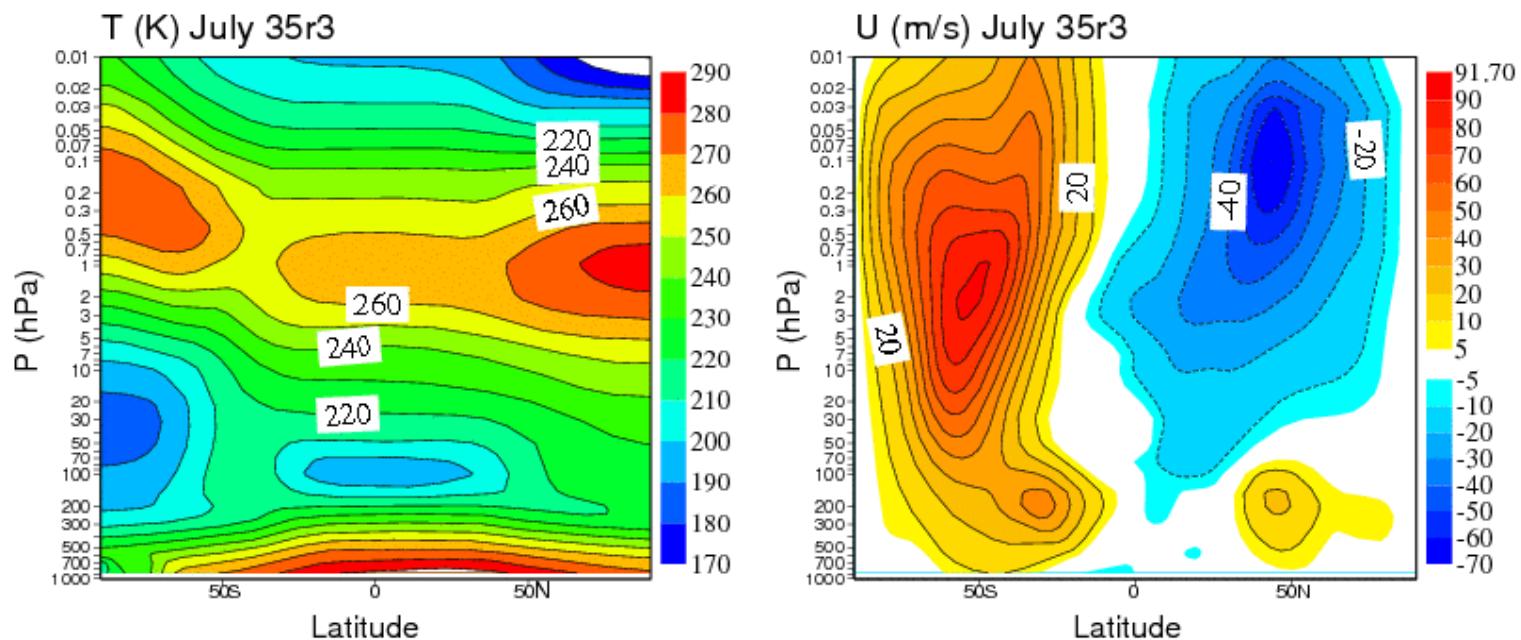
35r2

July climatology

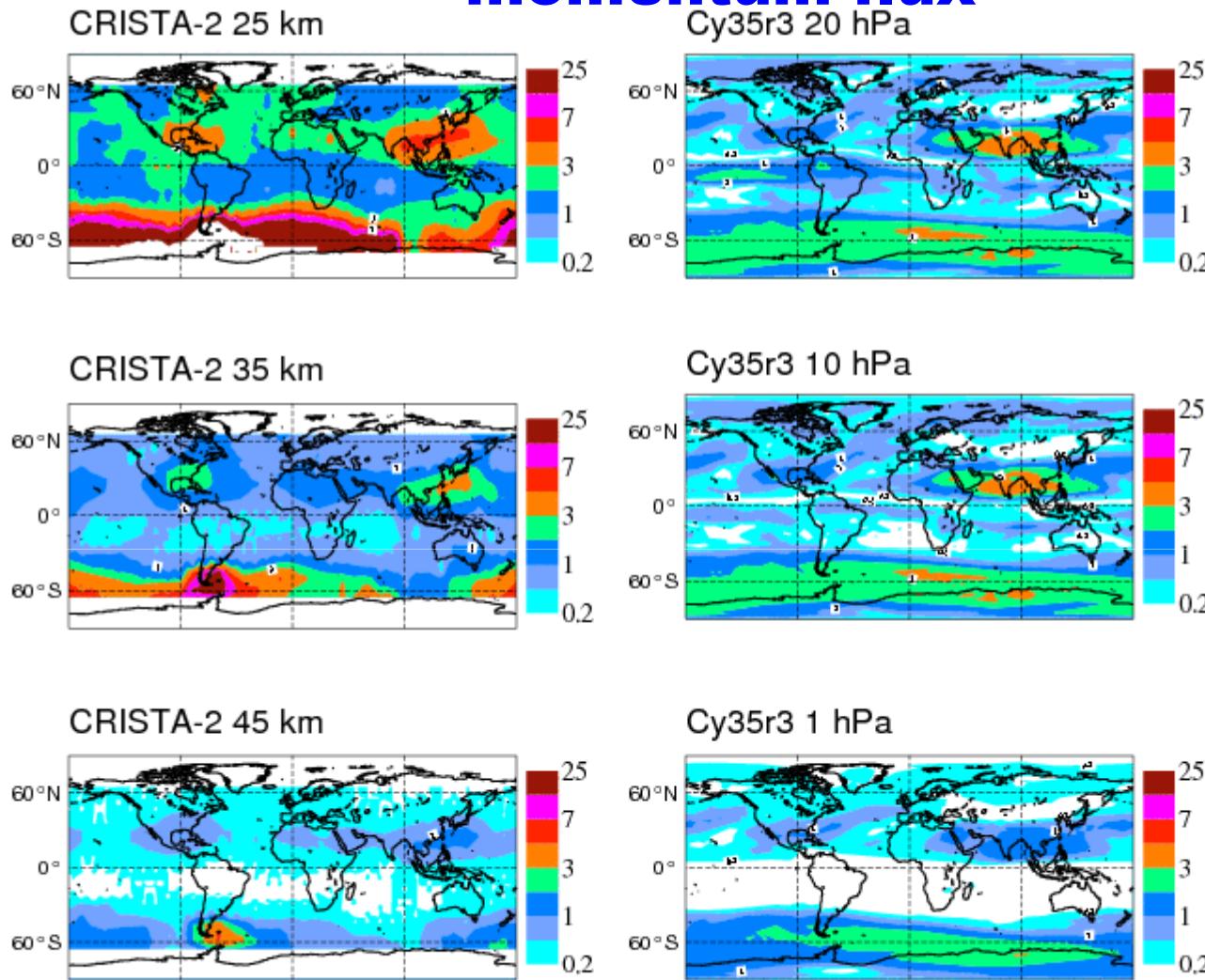
SPARC



35r3



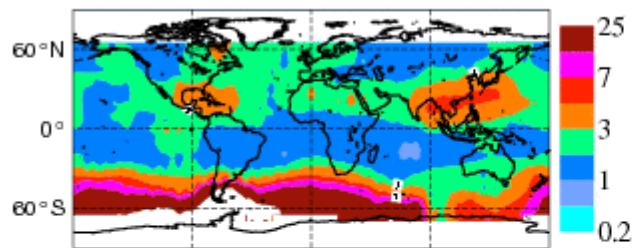
Parametrized non-orographic gravity wave momentum flux



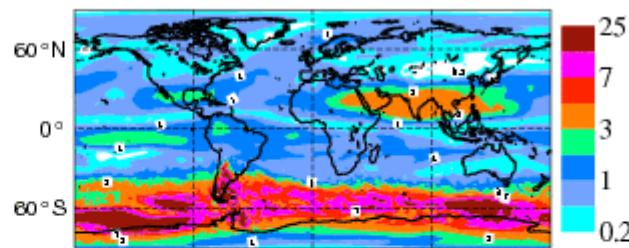
Comparison of observed and parametrized GW momentum flux for 8-14 August 1997 horizontal distributions of absolute values of momentum flux (mPa). Observed values are for CRISTA-2 (Ern et al. 2006). Observations measure temperature fluctuations with infrared spectrometer, momentum fluxes are derived via conversion formula.

Total = resolved + parametrized wave momentum flux

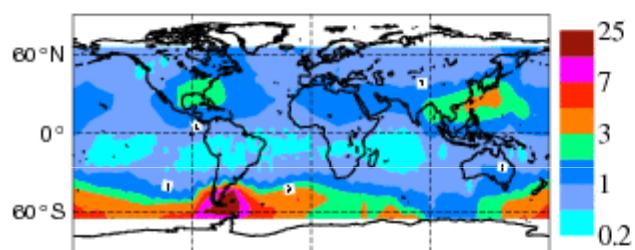
CRISTA-2 25 km



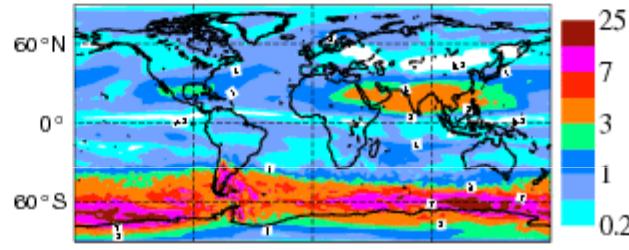
Cy35r3 20 hPa



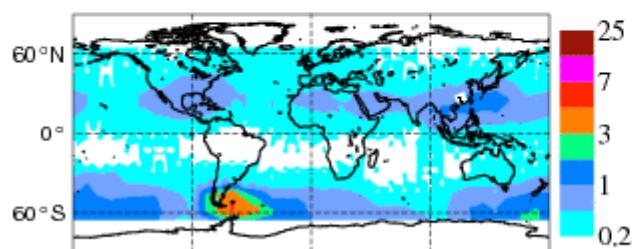
CRISTA-2 35 km



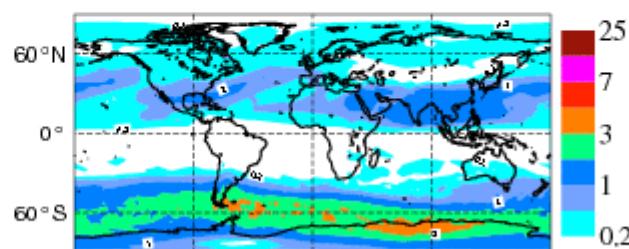
Cy35r3 10 hPa



CRISTA-2 45 km

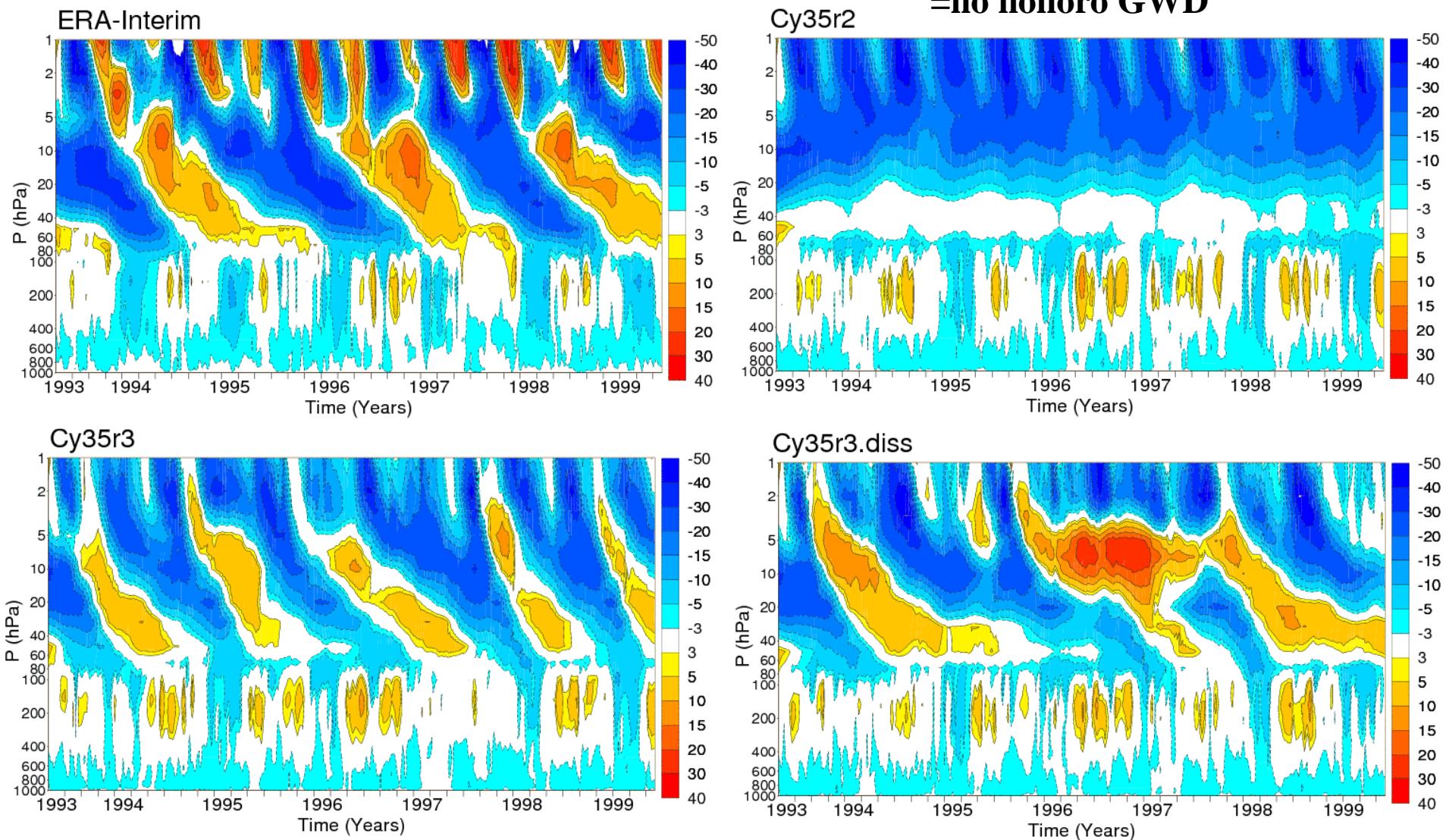


Cy35r3 1 hPa



A. Orr, P. Bechtold, J. Scinoccia, M. Ern, M. Janiskova (JAS 2009 to be submitted)

QBO : Hovmöller from free 6y integrations



ERA-Interim

- ◆ 20+ years: 1989-2009, continuing near-real time
 - ◆ Resolution: T255L60, 6-hourly (3-hourly for surface)
 - ◆ Forecast model version late 2006 (Cy31r2)
 - ◆ Analysis using 12-hourly 4D-Var
 - ◆ Variational bias correction of radiance data (VarBC)
-
- ◆ Member states: full access via MARS
 - ◆ Public web access via [ECMWF Data Server](#)
 - ◆ **Copy of complete archive at NCAR (<http://dss.ucar.edu>)**

Please visit: <http://www.ecmwf.int/research/era>

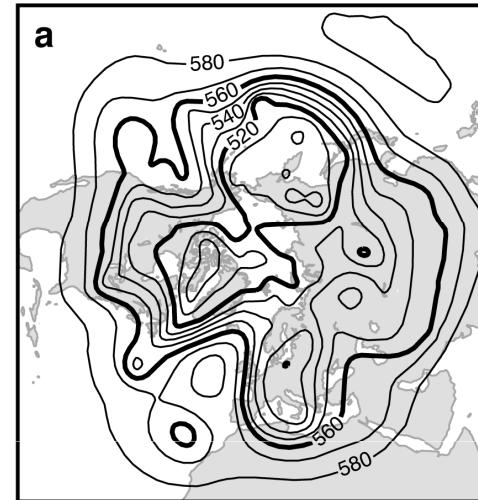
Atmospheric reanalysis at ECMWF

	ERA-15	ERA-40	ERA-Interim	ERA-75 (target)
TIME PERIOD	1979-1993	1957-2002	from 1989 onwards	from 1938 onwards
USERS	Meteorologists and Atmospheric Scientists Climate Scientists and Wider Earth Science Community Additional "Environmental Users"			European Stakeholders GMES Core & Downstream services
INPUT DATA ACCESS	Mixed Observational Data Formats in Archive Observation Quality Feedback Information			Unified, Consolidated Database Facility Internet Access
GRIDDED PRODUCTS	Model Fields (GRIB format)			Real-time Product Database Essential Climate Variables Internet Access
ATMOSPHERE	Assimilation OI 31 levels 150km	Assimilation 3DVAR 60 levels 125km	Assimilation 4DVAR 60 levels 80km	Assimilation weak-constraint 4DVAR 91 levels 40 km Improved Observations
LAND	Forcing	Model	Improved Model	Improved Model & Assimilation Coupling
OCEAN & SEA-ICE	SST/ice Forcing		Improved SST/ice Forcing Wave Model	Improved SST/ice Coupling
CHEMISTRY		Forcing	Improved Forcing	Improved Interaction
IMPACT	Enhance Understanding of Atmospheric Variability, Leading to Improved Models Investigate Past Weather and Climate, Assess Observing System Impact		Monitor Near Real-time Climate with Traceability to Input Data	Facilitate Environmental Decisions, Enable New Applications of GMES, Assess Regional Climate Change & Risks via Regional Reanalyses, Improve Earth System Modeling, Maximize Benefits from Earth Observation Infrastructure

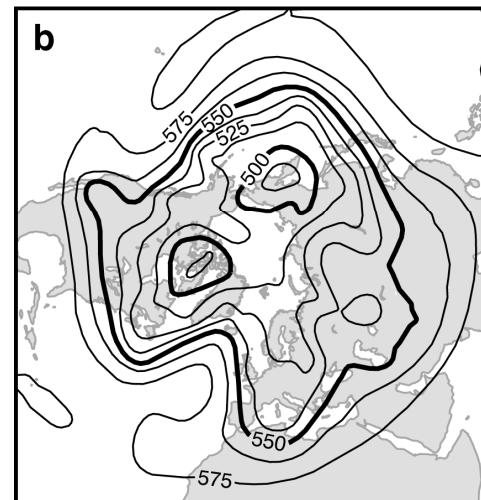
Reanalysis of sparse observations

4D-Var CONTROL
All observations

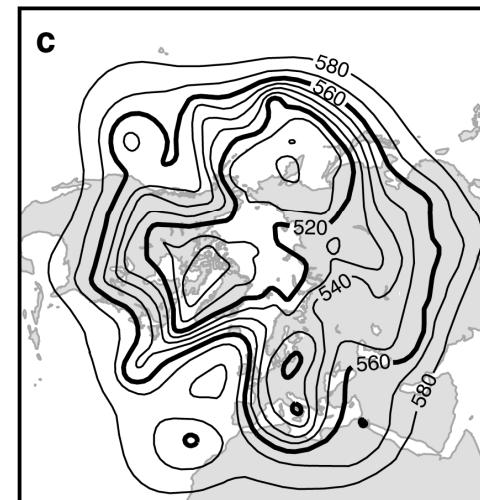
(15 February 2005 00 UTC)



3D-Var
“Surface
pressure
observations
only”



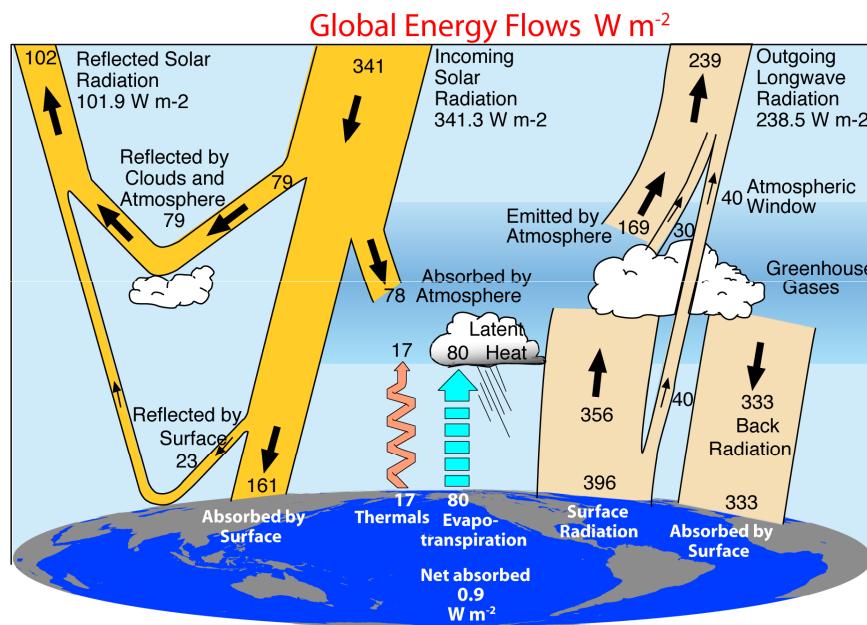
4D-Var
“Surface
pressure
observations
only”



Whitaker,
Compo, and
Thépaut 2009

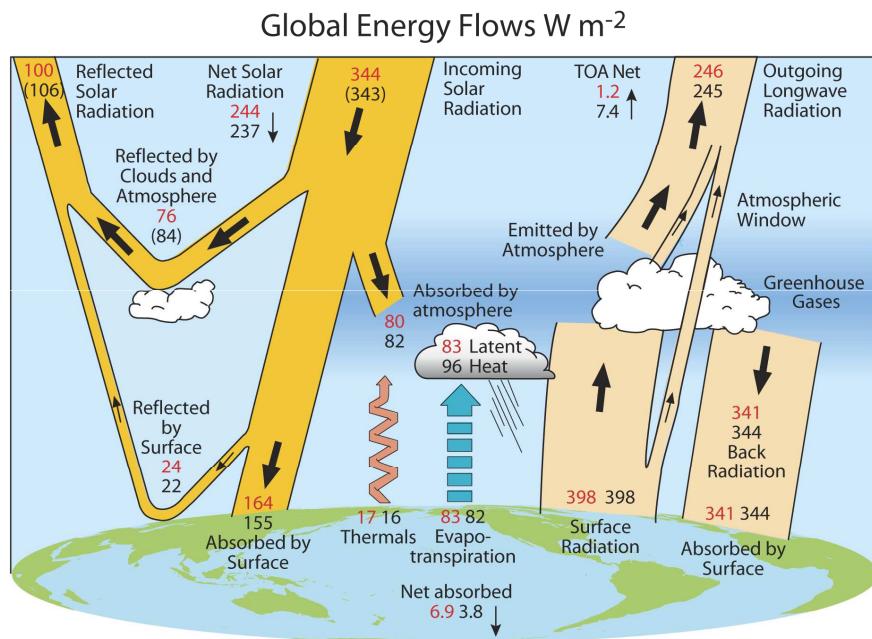
Energy budget

Trenberth et al. 2009:
Earth's global energy budget



ERA-Interim
1989-2008

ERA-40
1989-2001



- TOA balance improved in ERA-Interim
- Surface energy balance worse, esp. over oceans

Mean age of air in the lower stratosphere

Based on 20-year CTM runs,
using reanalysed winds
from ERA-40 and ERA-
Interim

Observational estimates
derived from ER-2
aircraft measurements of
CO₂ and SF₆

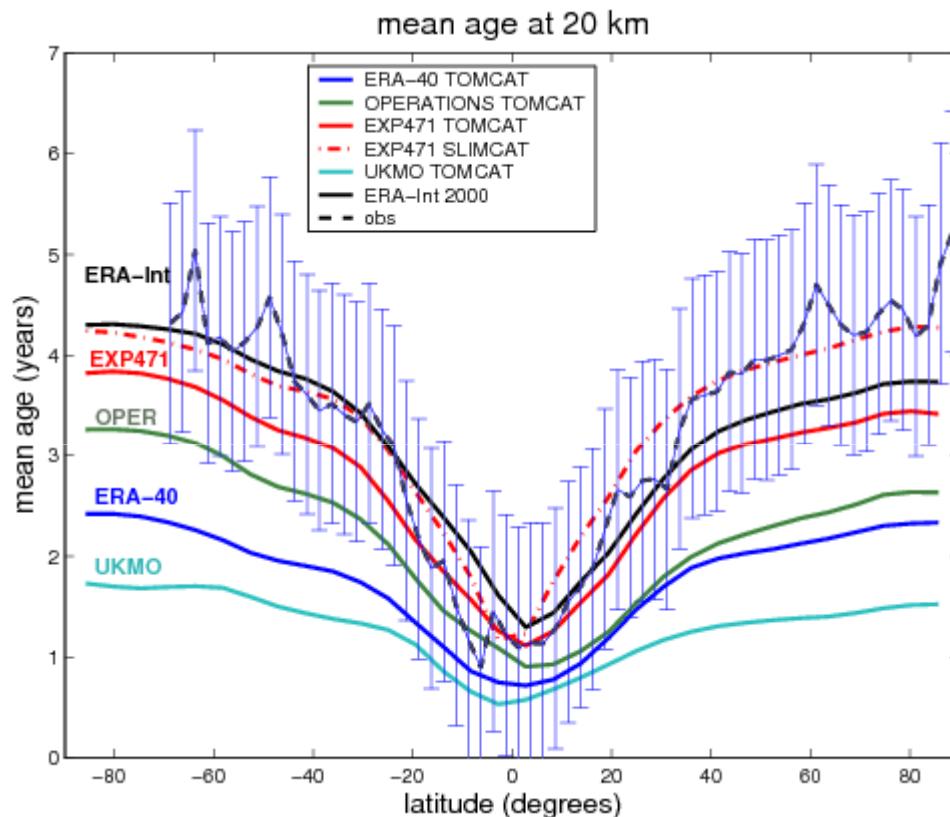
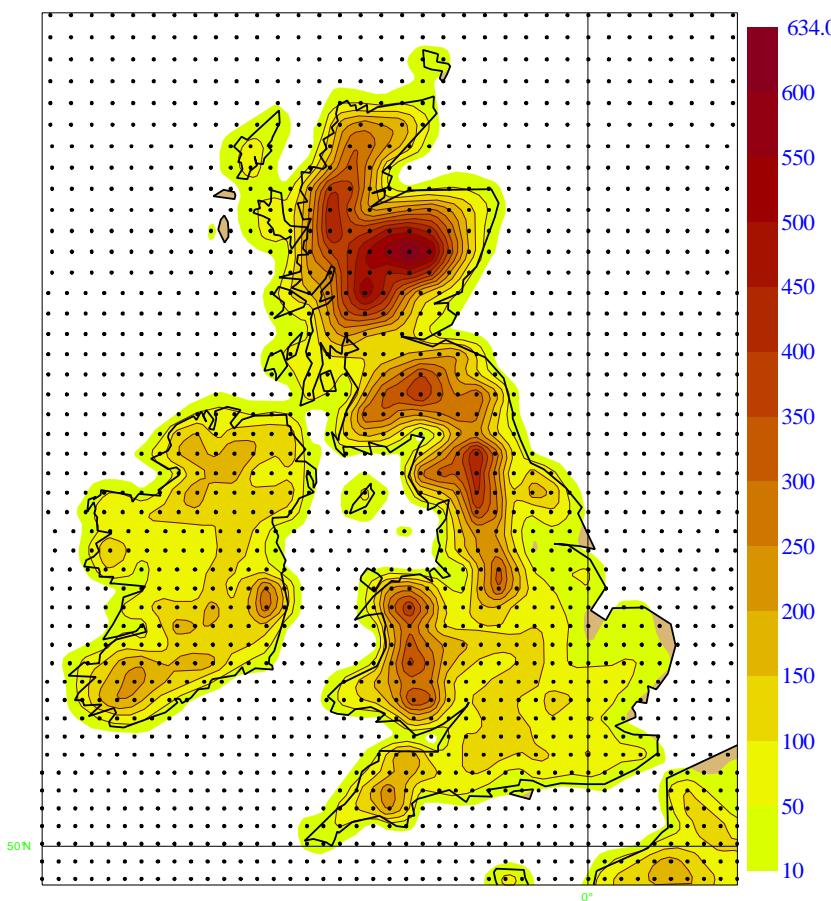


Figure updated from Monge-Sanz et al. 2007

High resolution T_L1279

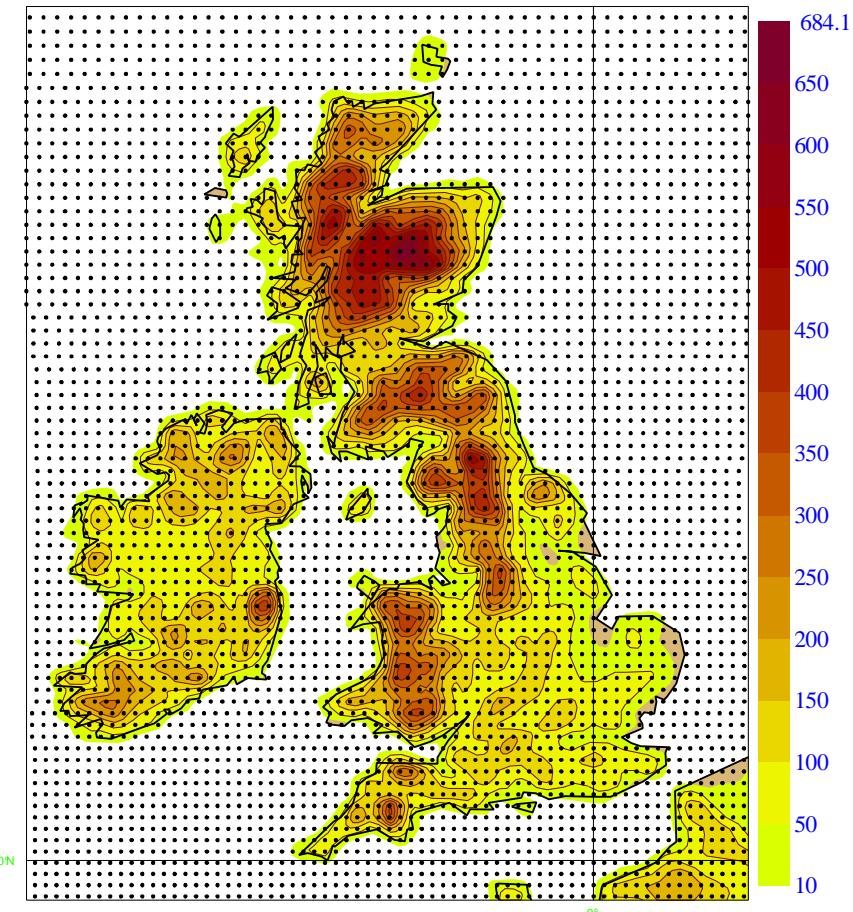
- ◆ 4D-Var configuration tested extensively
 - T1279 L91 outer loop (time-step 600s)
 - T159/T255/T255 inner loops (time-steps 1800s)
- ◆ Ongoing validation, so far 267 days of assimilation and forecast with CY35R3

T799



25 km grid-spacing
(843,490 grid-points)

T1279



16 km grid-spacing
(2,140,704 grid-points)

(thanks to Mariano Hortal who actually created these grids)

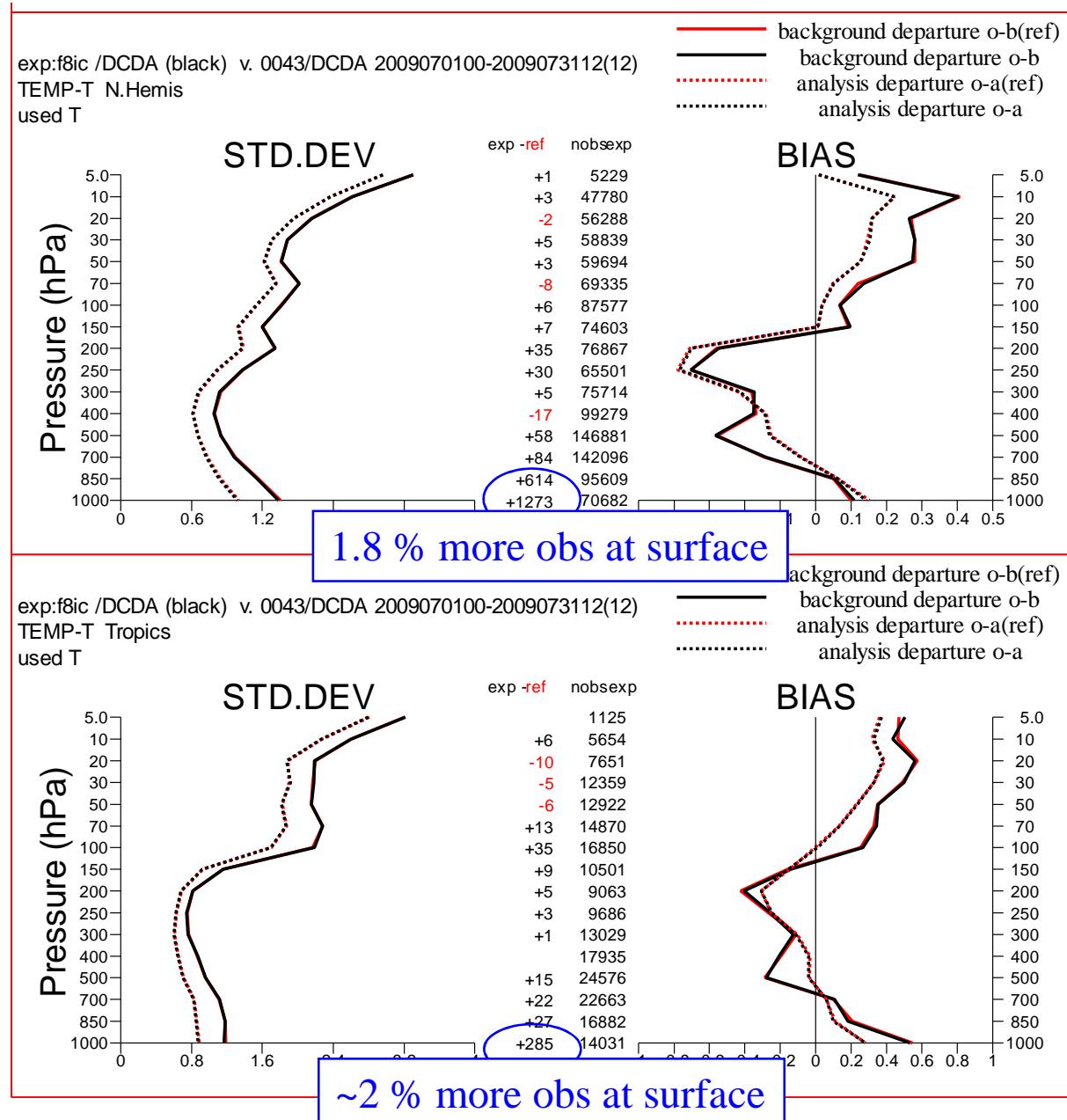
Fit to TEMP T observations (average for July 2009)

f8ic T1279
(black)

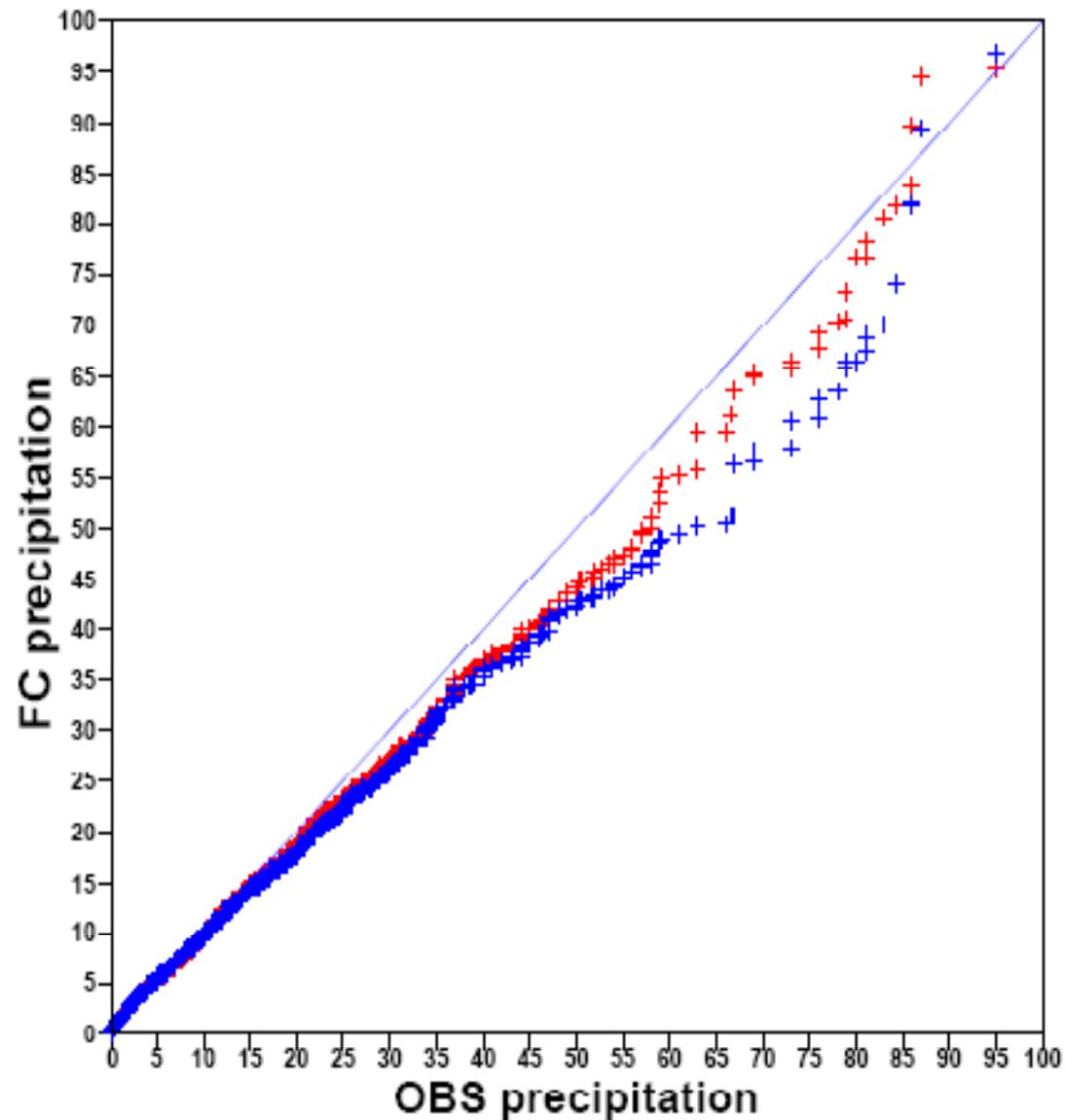
Control (red)
35r3 e-suite
0043

NH

Tropics



T1279 & T799 versus Obs precipitation Europe



T1279

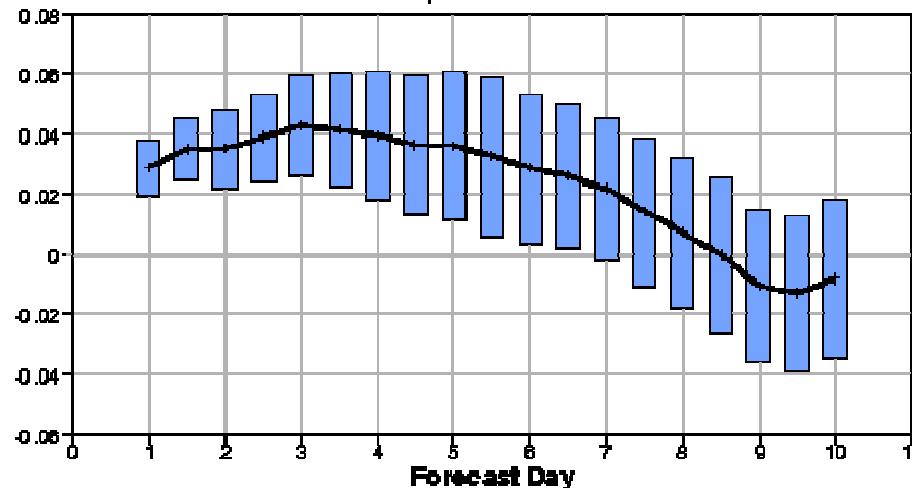
T799

from 36h
forecast
winter

High resolution
improves on high
precipitation
amounts

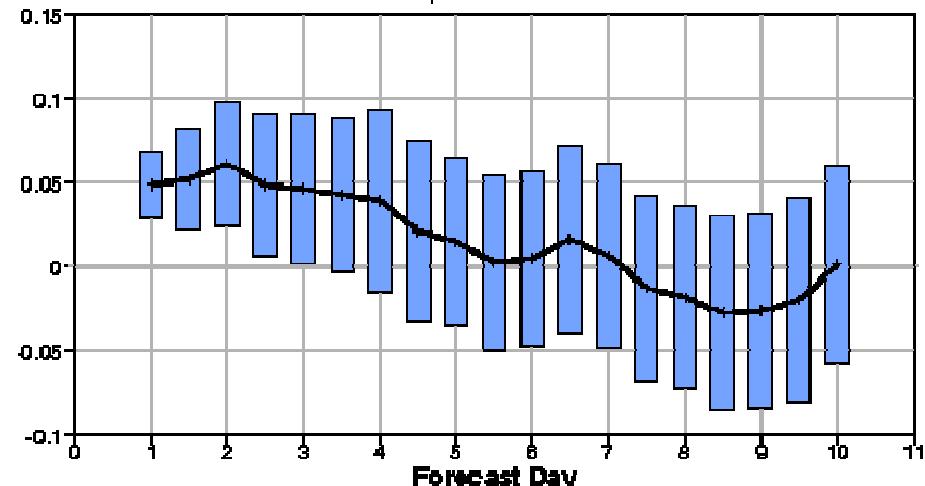
NH

control normalised f7zb minus f7z1
Anomaly correlation forecast
Nhem Lat 20.0 to 90.0 Lon -180.0 to 180.0
Date: 20081201 00UTC to 20081231 00UTC
500hPa Geopotential 00UTC
Confidence: 95%
Population: 267

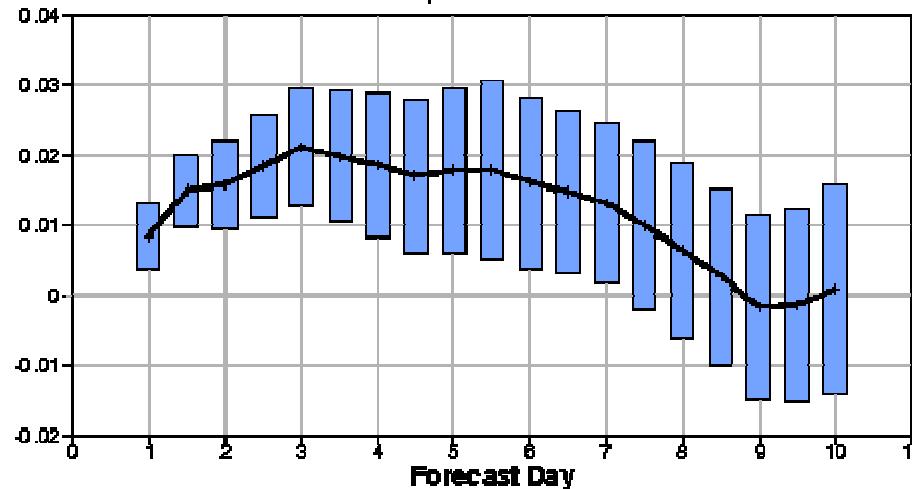


Europe

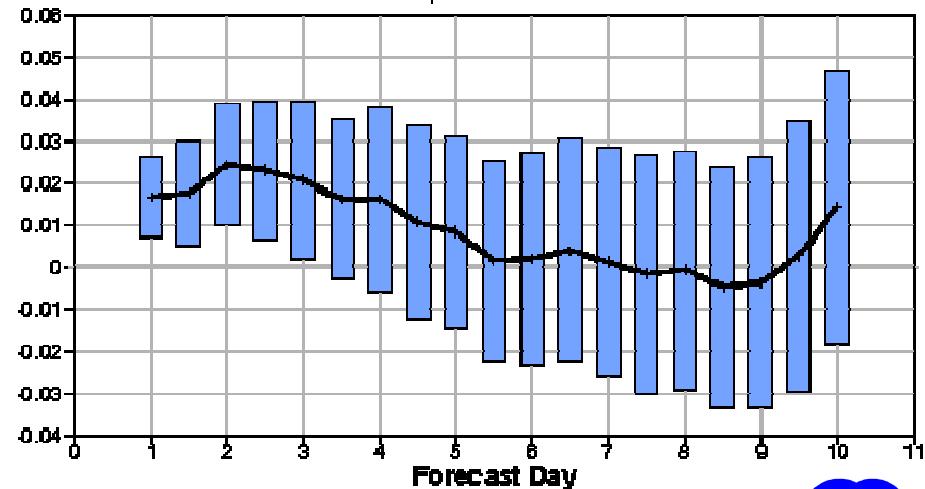
control normalised f7zb minus f7z1
Anomaly correlation forecast
Europe Lat 35.0 to 75.0 Lon -12.5 to 42.5
Date: 20081201 00UTC to 20081231 00UTC
500hPa Geopotential 00UTC
Confidence: 95%
Population: 267



control normalised f7z1 minus f7zb
Root mean square error forecast
Nhem Lat 20.0 to 90.0 Lon -180.0 to 180.0
Date: 20081201 00UTC to 20081231 00UTC
500hPa Geopotential 00UTC
Confidence: 95%
Population: 267



control normalised f7z1 minus f7zb
Root mean square error forecast
Europe Lat 35.0 to 75.0 Lon -12.5 to 42.5
Date: 20081201 00UTC to 20081231 00UTC
500hPa Geopotential 00UTC
Confidence: 95%
Population: 267



The nonhydrostatic IFS

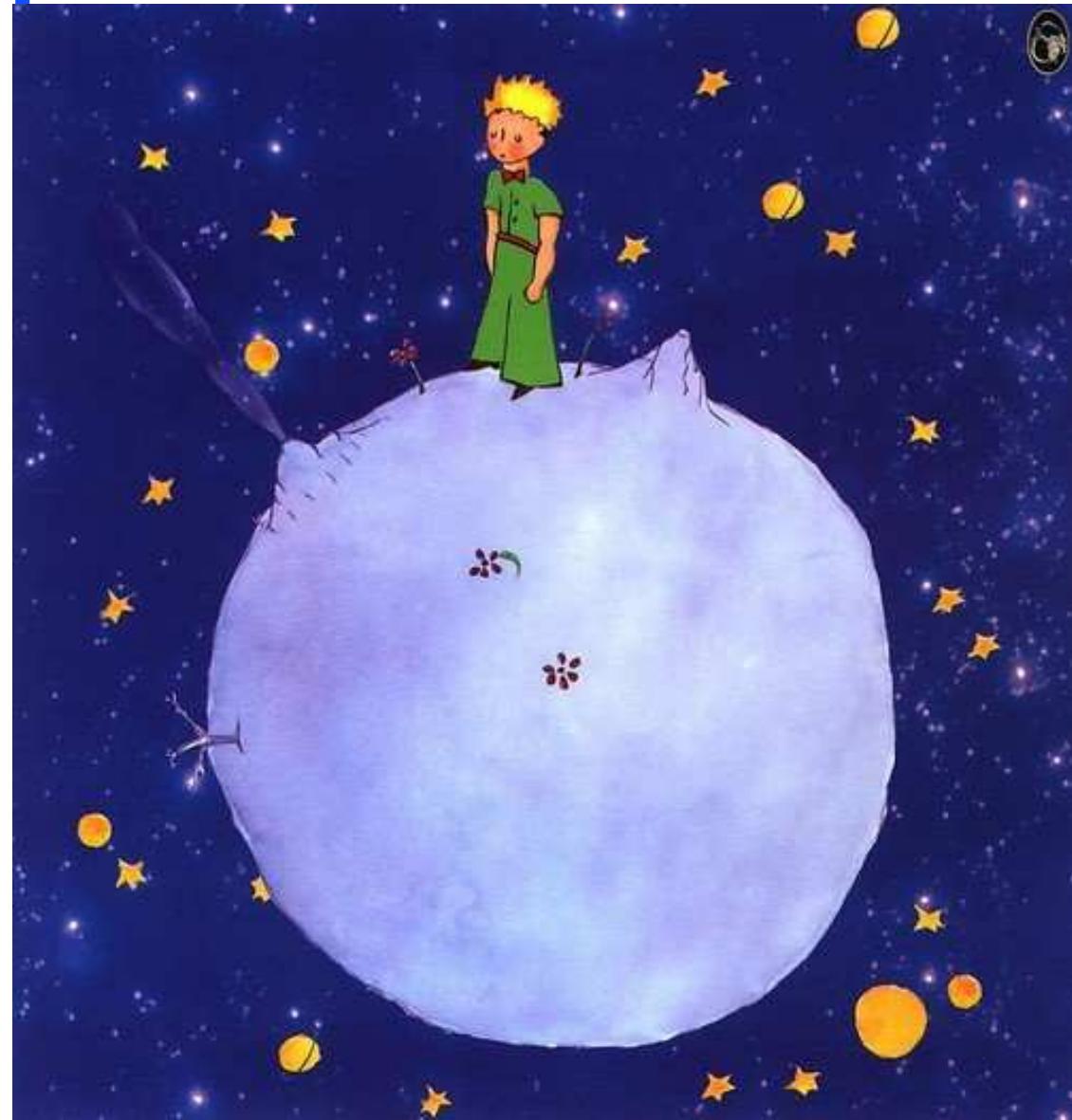
- ◆ Developed by Météo-France and its ALADIN partners
Bubnová et al., (1995); ALADIN (1997); Bénard et al. (2004,2005,2009)
- ◆ Made available in IFS/Arpège by Météo-France (*Yessad, 2008*)
- ◆ Testing of NH-IFS described in Techmemo TM594 (*Wedi et al. 2009*)

Hierarchy of test cases

- ◆ *Shallow water*
- ◆ **Acoustic waves**
- ◆ **Gravity waves**
- ◆ **Planetary waves**
- ◆ **Convective motion**
- ◆ **Idealized dry atmospheric variability and mean states**
- ◆ *Moist simulations and dynamics-physics interaction*
- ◆ **Seasonal climate, intraseasonal variability**
- ◆ **Medium-range forecast performance at hydrostatic scales**
- ◆ *High-resolution forecasts at nonhydrostatic scales*

Local- and synoptic-scale simulations on the sphere ...

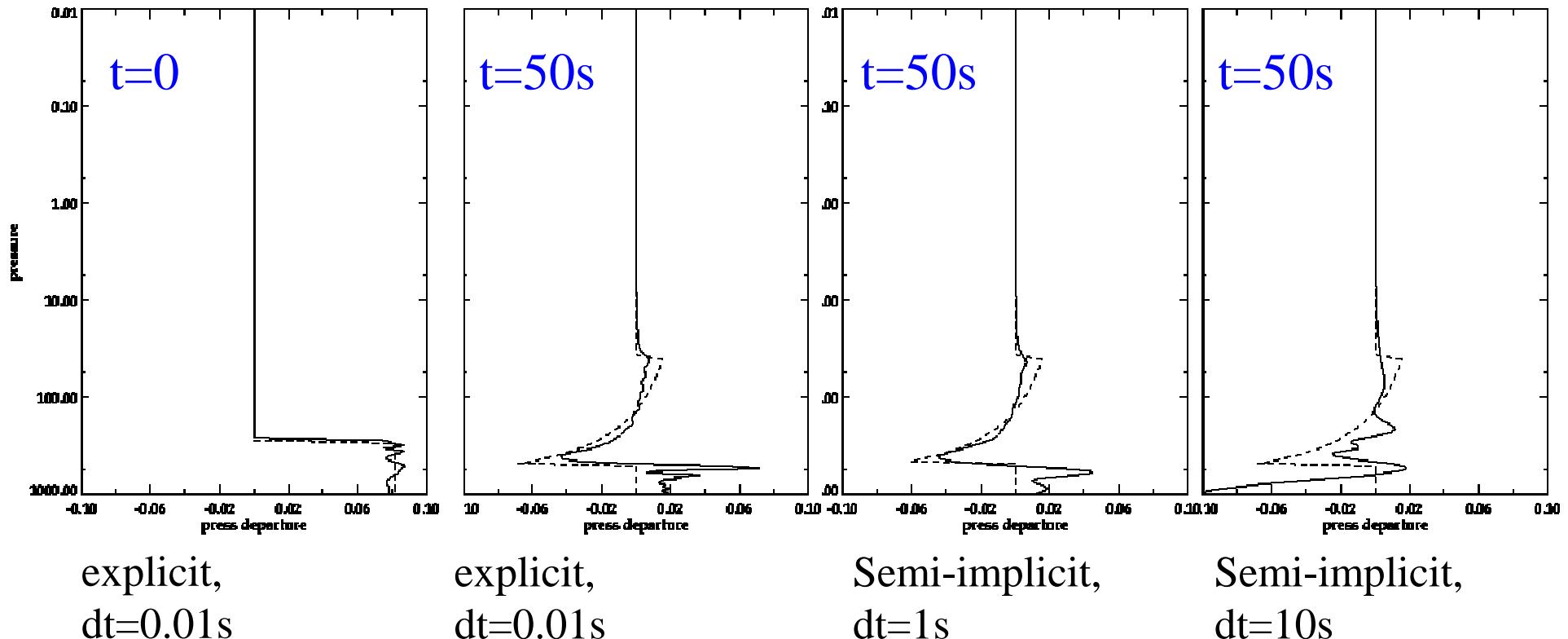
The size of the computational domain is reduced without changing the depth or the vertical structure of the atmosphere by changing the radius ($a < a_{\text{Earth}}$)



Wedi and Smolarkiewicz, Q. J. R. Meteorol. Soc. 135: 469-484 (2009)

A spherical acoustic wave in a stratified (isothermal) atmosphere

analytical solution in dashed line



explicit,
 $\Delta t=0.01s$

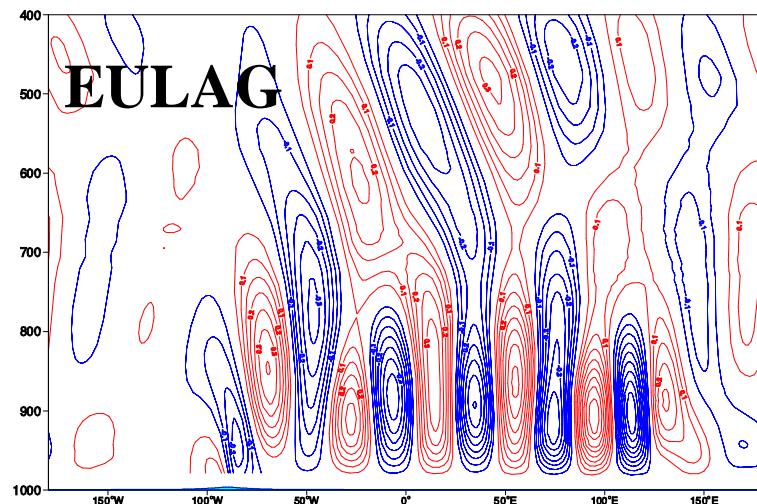
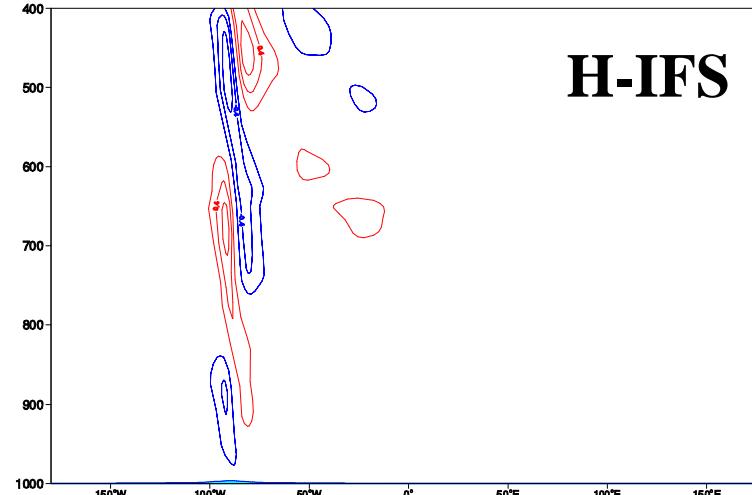
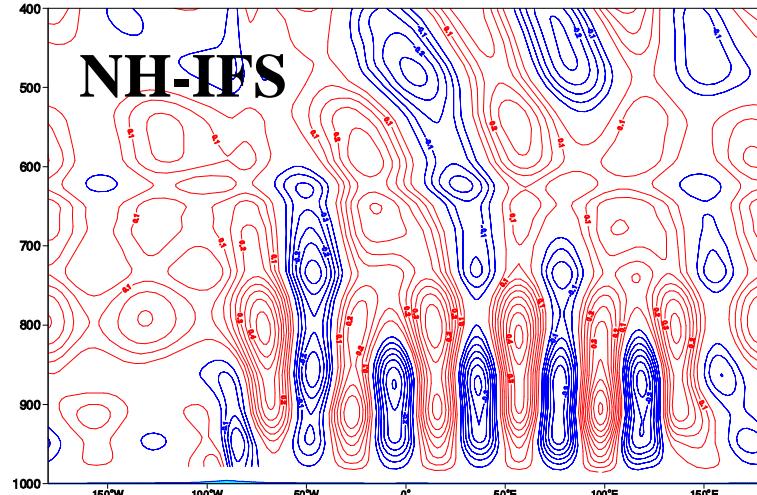
explicit,
 $\Delta t=0.01s$

Semi-implicit,
 $\Delta t=1s$

Semi-implicit,
 $\Delta t=10s$

NH-IFS T_L159L91

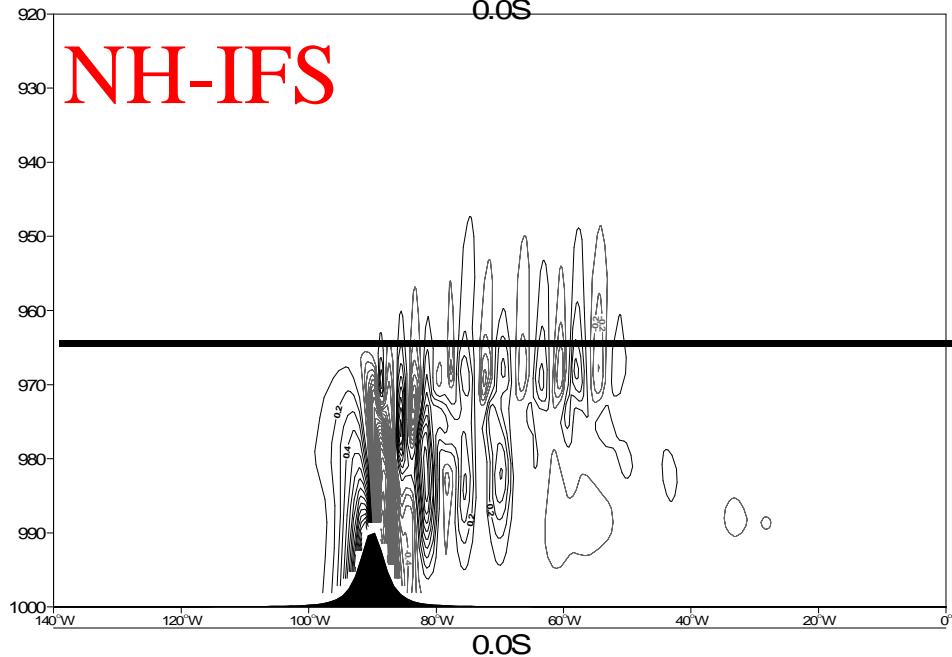
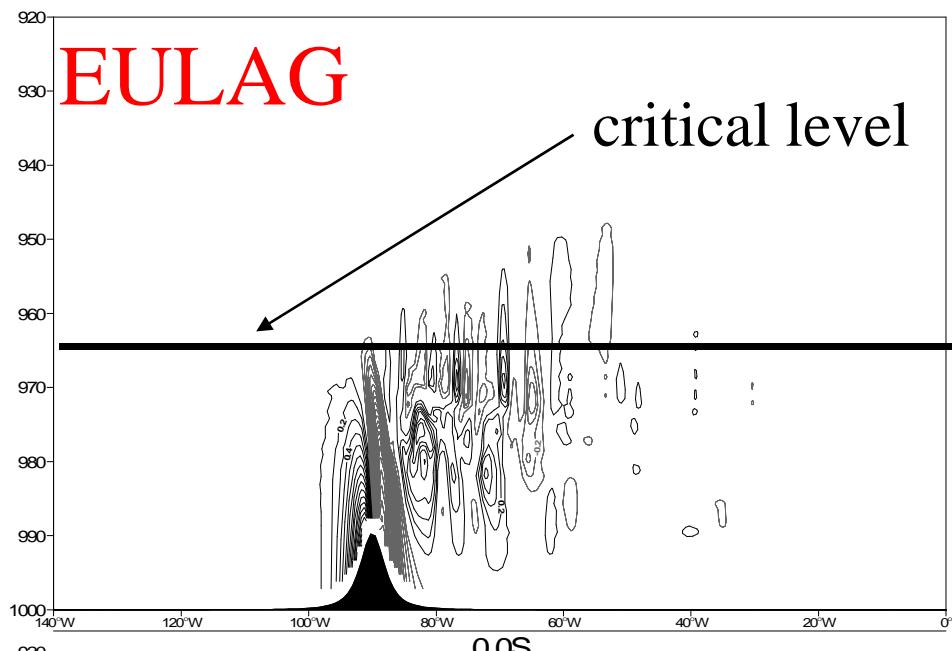
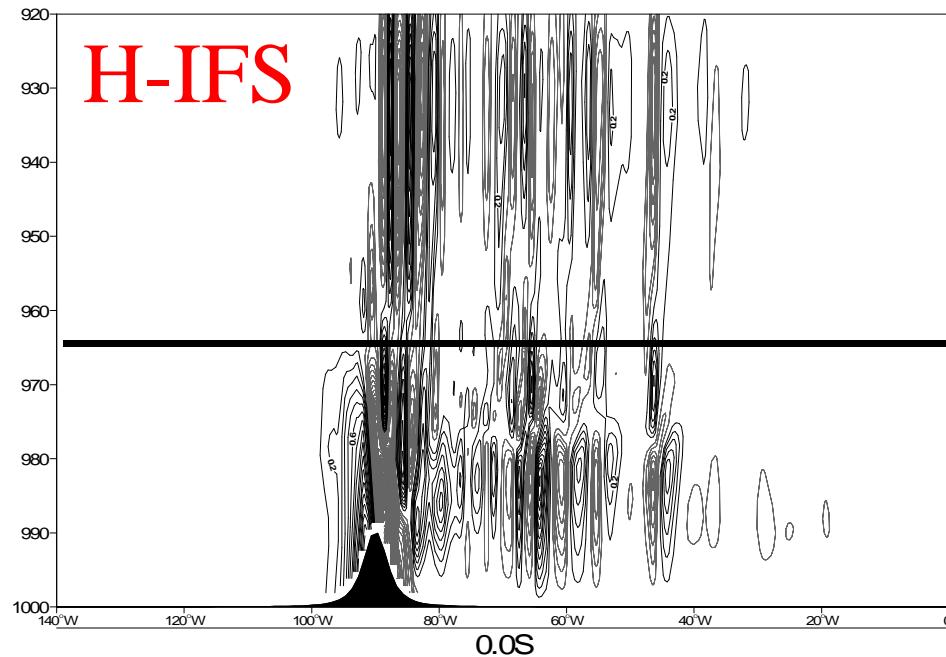
Quasi two-dimensional orographic flow with linear vertical shear



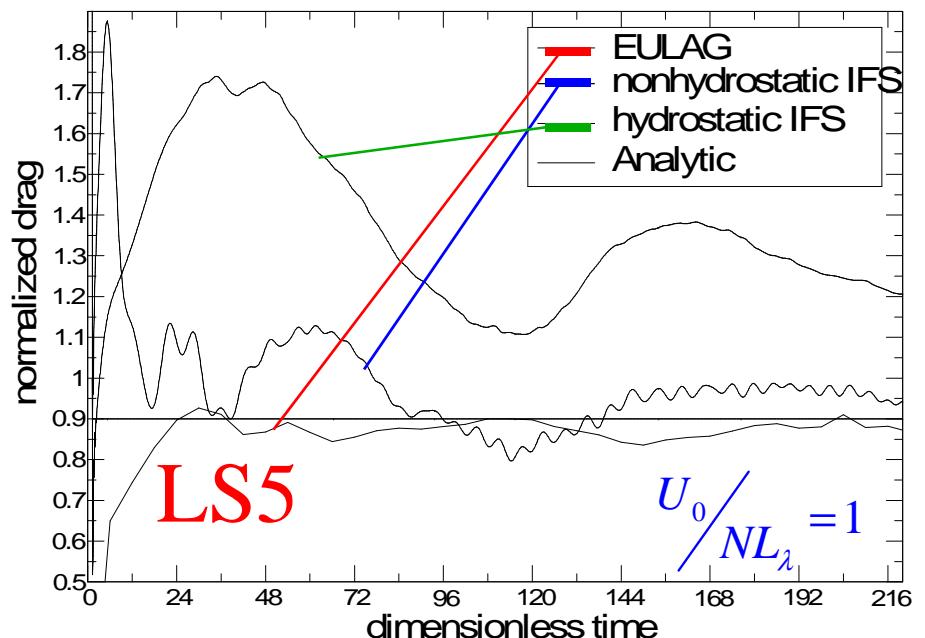
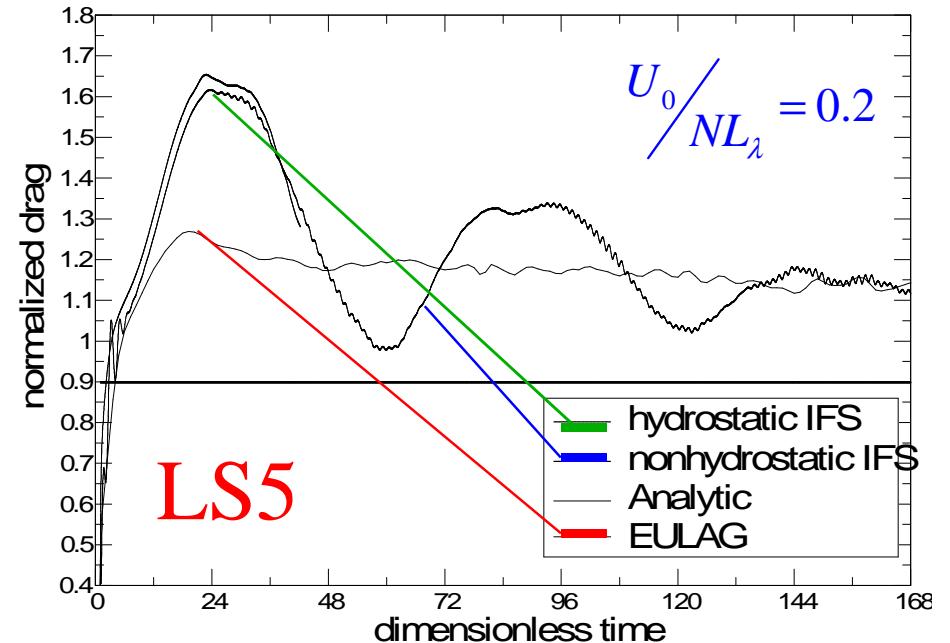
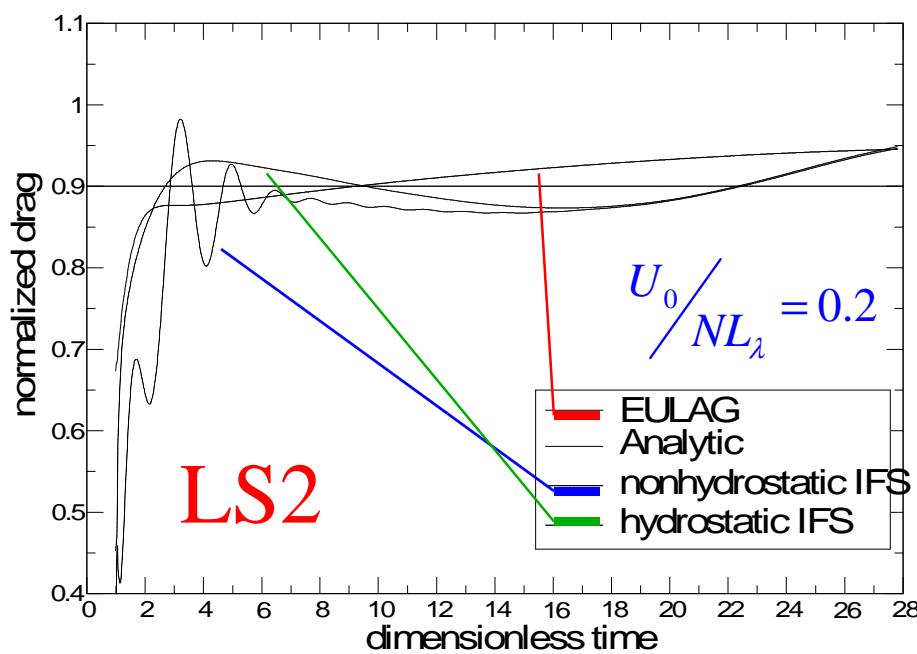
The figures illustrate the correct horizontal (NH) and the (incorrect) vertical (H) propagation of gravity waves in this case (Keller, 1994). Shown is vertical velocity.

The critical level effect on linear and non-linear flow past a three-dimensional hill

LS5 $\frac{U_0}{NL_\lambda} = 1$



Mountain drag

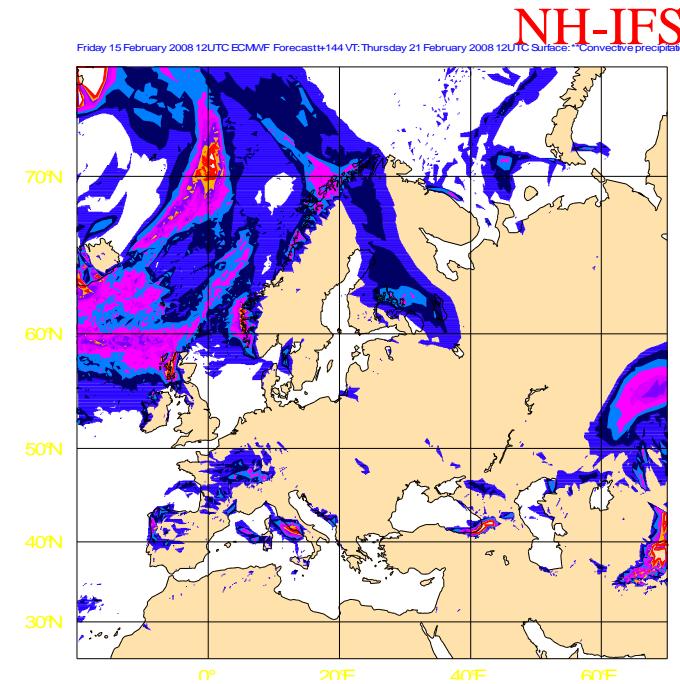
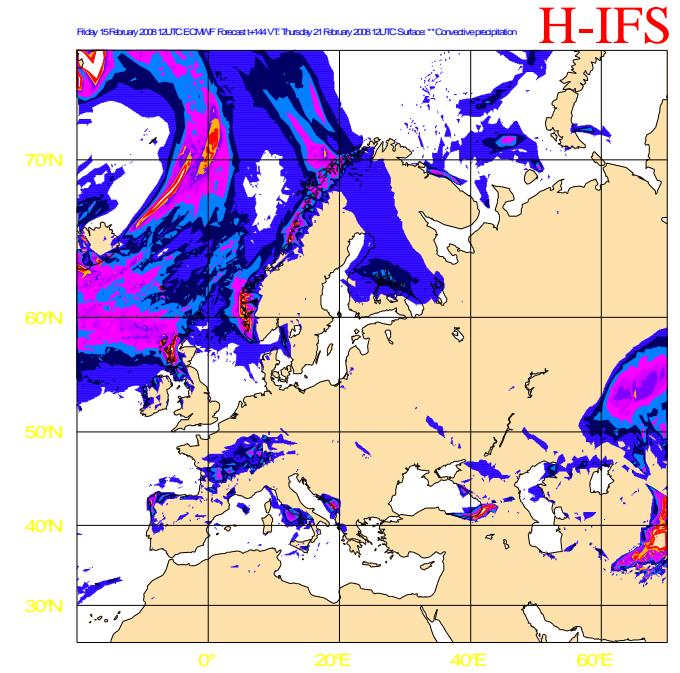
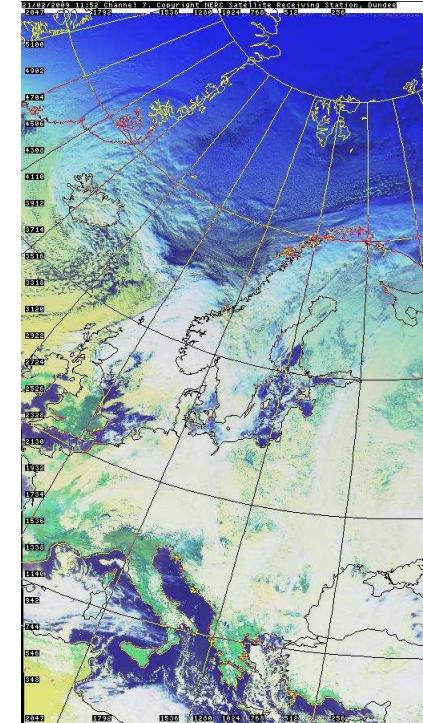
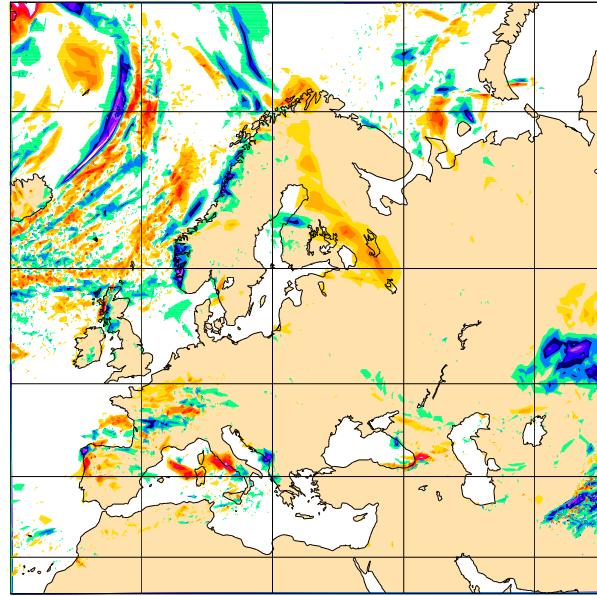


Medium-range forecast performance at hydrostatic scales

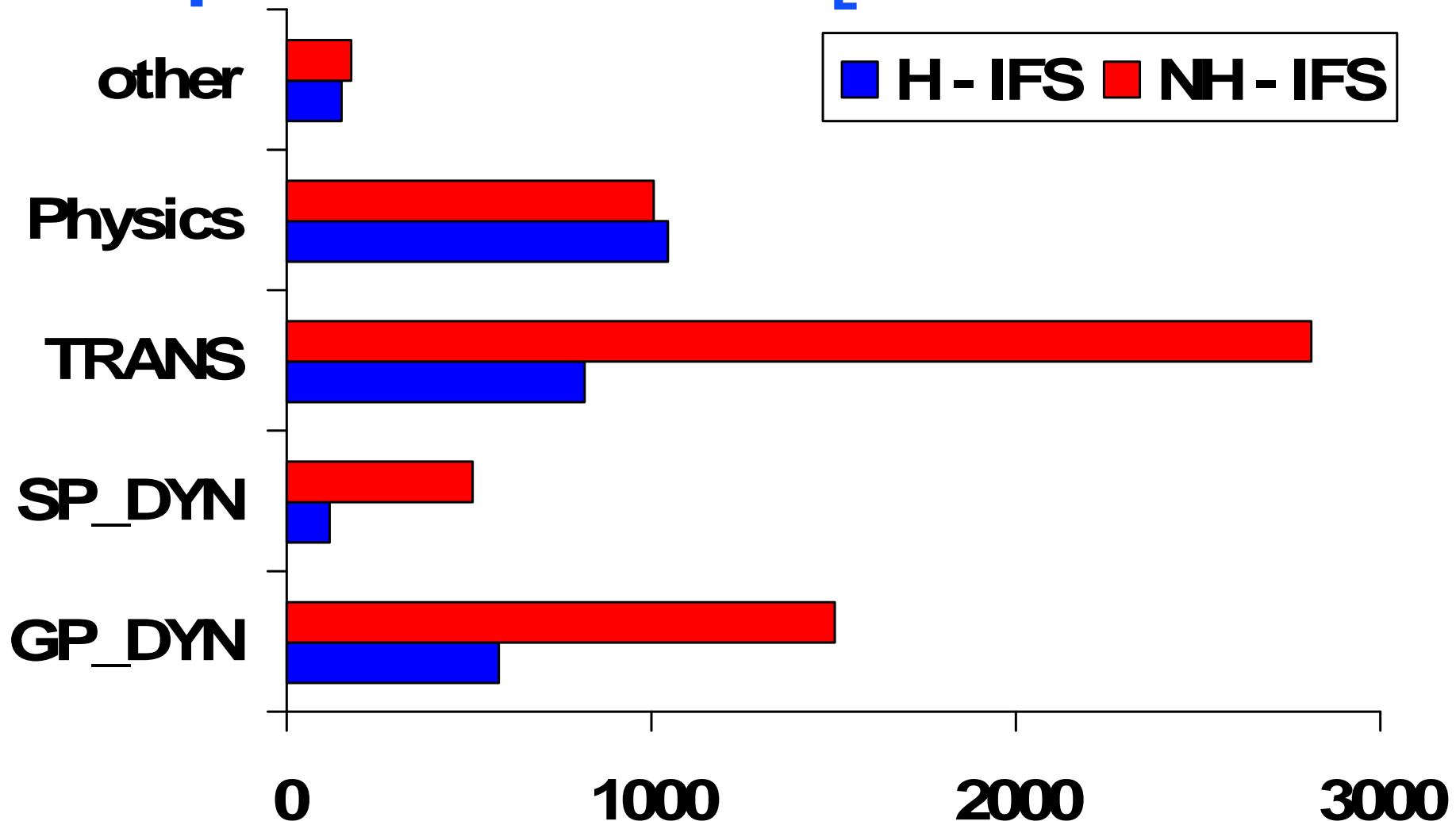
- ◆ Tropospheric scores are the same T799L91 (41 cases) 33r2, T1279 (49 cases) 35r1; Stratospheric scores the same with LVERTFE=T (here the NH-IFS model uses the linear system and the semi-implicit calculations in FD! -> Karim Yessad)
- ◆ So far physics coupling to the dynamics in the ICI scheme in final iteration only and using the hydrostatic physics “as is”. Open issues with physics-dynamics coupling towards cloud-resolving simulations!

T_L 2047 ~ 10 km grid-size

Total precipitation forecast difference
at $t_0 + 6$ days

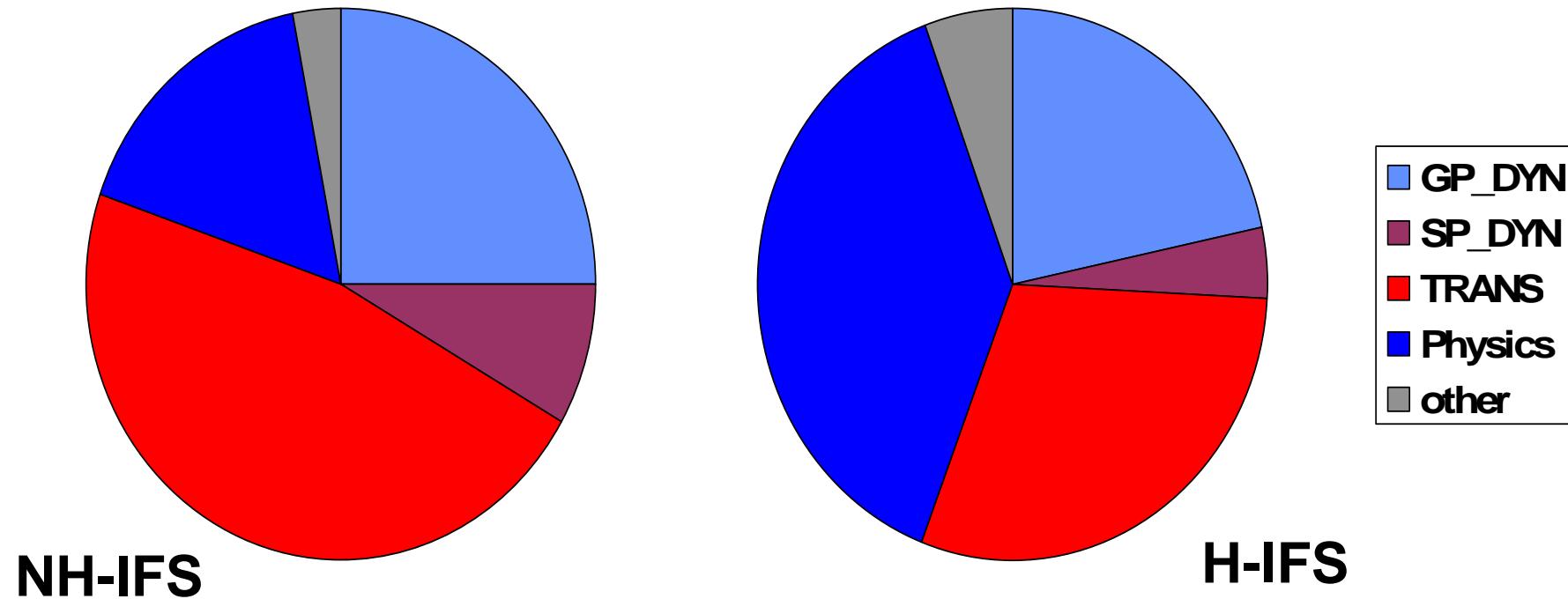


Computational Cost at $T_L 2047$



Total cost increase 106 %

Computational Cost at $T_L 2047$



NH-IFS status summary

- ◆ Works reasonably well for various idealized test cases
- ◆ Deep and shallow atmosphere options available
- ◆ Seasonal climate and scores almost identical to hydrostatic model at hydrostatic scales
- ◆ Little or no benefit found so far at T_{L2047}
- ◆ Computational cost 2 x at T_{L2047} is an issue !
 - ◆ Renewed interest in “Fast Legendre Transform” (*Tygert, 2008*)

Additional slides

Legendre transforms

- ◆ Changes to transform package went into cycle 35r3 that allow the computation of Legendre functions and Gaussian latitudes in double precision following (*Schwarztrauber, 2002*) (-> Mats Hamrud).
- ◆ **Note:** the increased accuracy 10^{-13} instead of 10^{-12} in the “Courtier and Naughton procedure” leads to slightly more points near the pole for all resolutions (in addition all resolutions have an odd feature of artificially excluding rows with 625 and 1250 points).

Fast Legendre transform

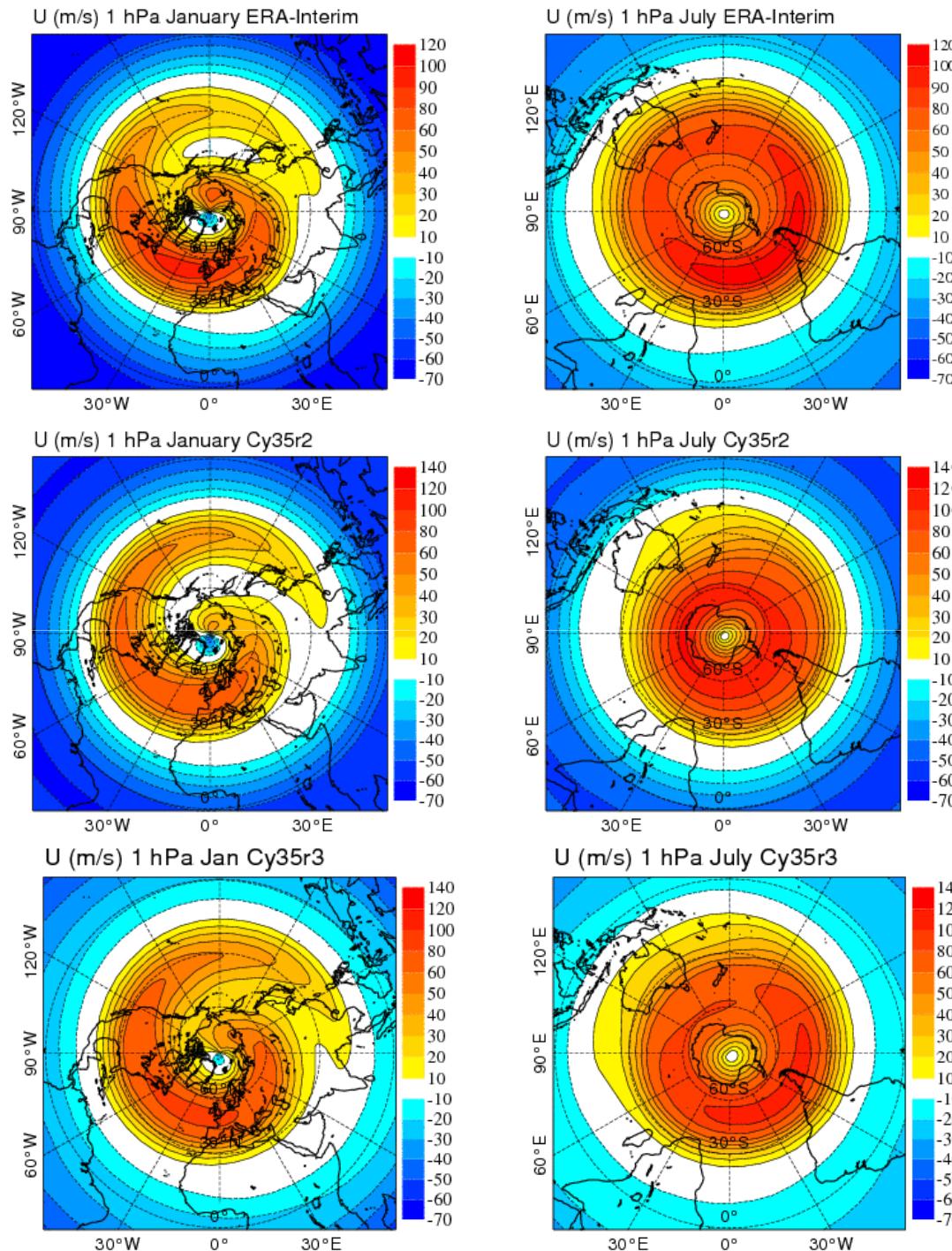
- ◆ Some time has been spent to start coding an *FMM (fast multipole method) based Legendre transform* following (*Tygert, JCP, 2008*). (-> Mats Hamrud and Mike Fisher).
- ◆ 1) A recursive *Cuppen divide-and-conquer algorithm* computes the Eigenvalues and Eigenvectors of a tridiagonal matrix and (“as-a-by-product”) it allows to apply a square matrix of normalized Eigenvectors to any arbitrary vector (which is what we really want to do).
- ◆ 2) The arising *matrix-vector multiply* can be accelerated by using a *FMM method* as the Eigenvector elements have a special form $\sim z/(d-\lambda)$, that is one can identify different levels/groups of interaction as a function of distance to the Eigenvalue λ .
- ◆ The devil is in the detail ...

January
NH

ERA

Cy35r2
Operational
since March
2009

Cy35r3
Operational
in summer
2009 with
GWD + GHE

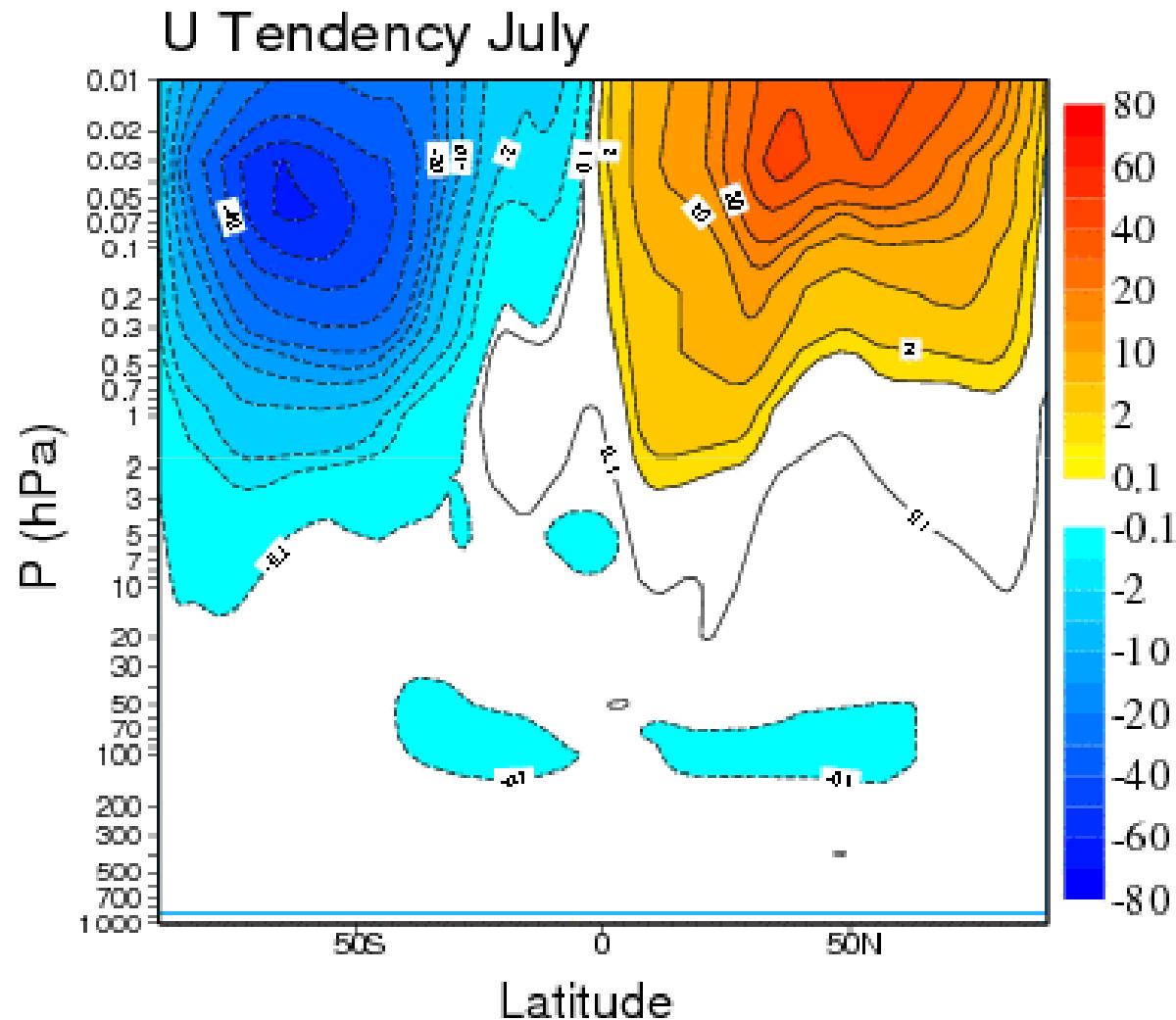


July
SH

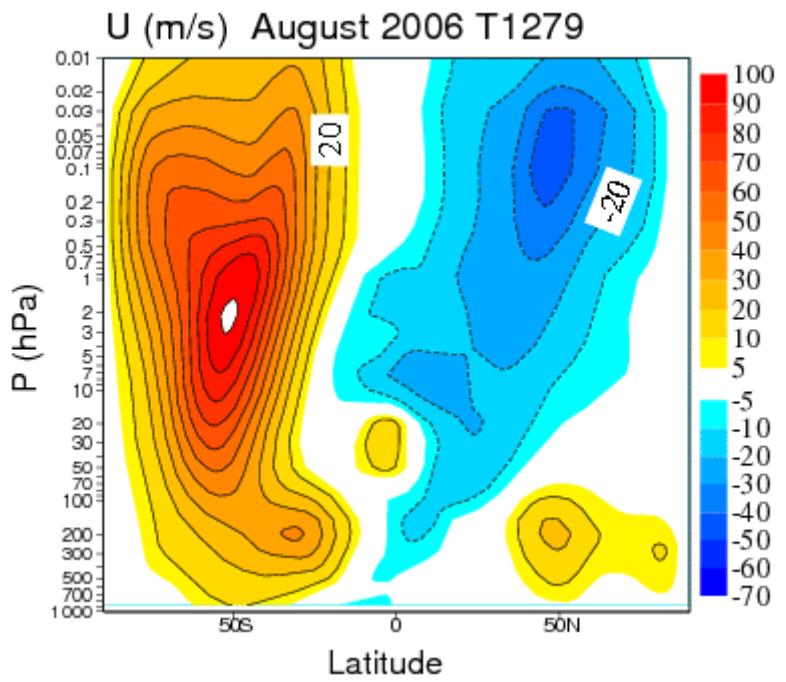
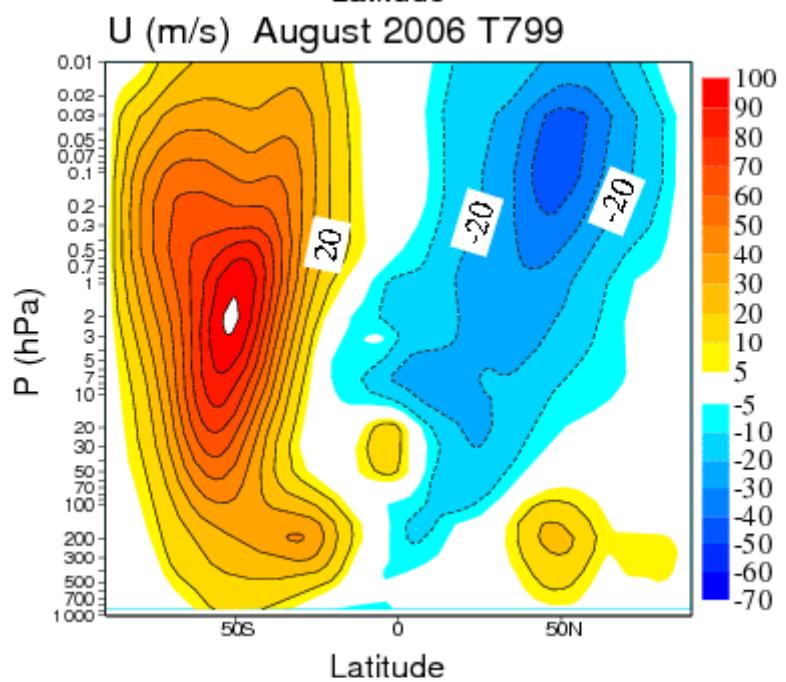
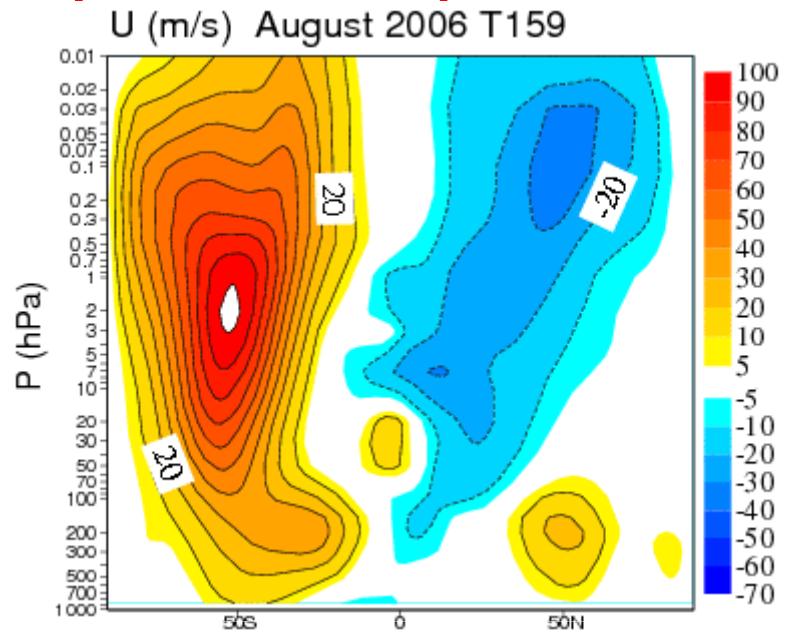
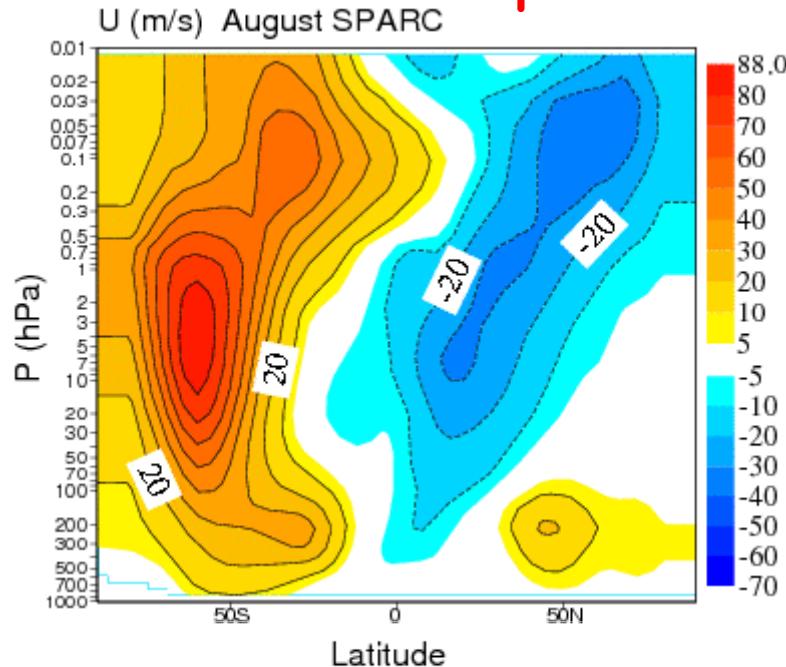
Polar
winter
vortex

SH wintertime
vortex is quasi-
symmetric, but
not NH polar
vortex, due to
braking quasi-
stationary Rossby
waves emanating
in the
troposphere

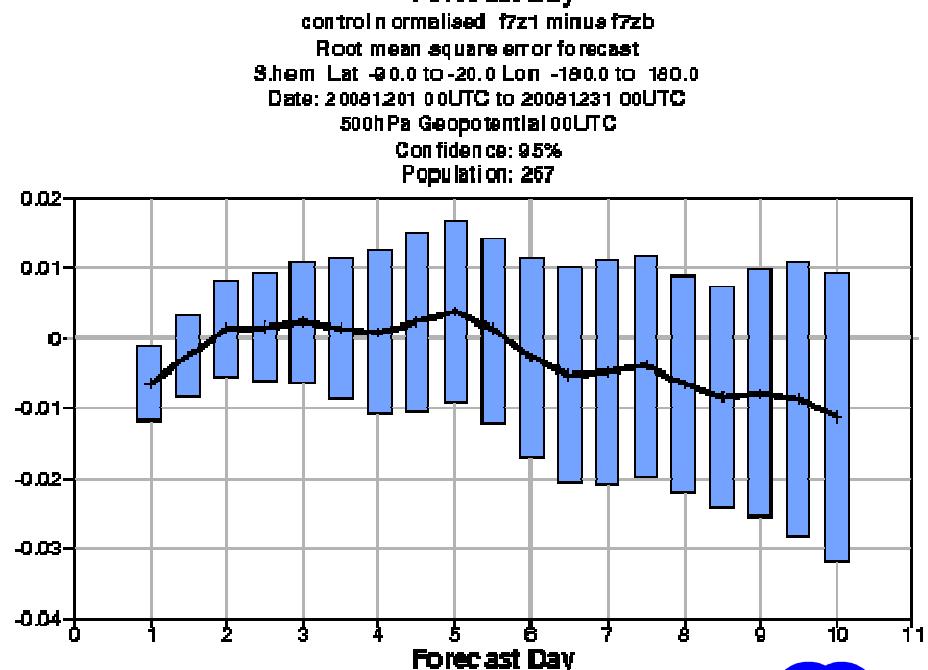
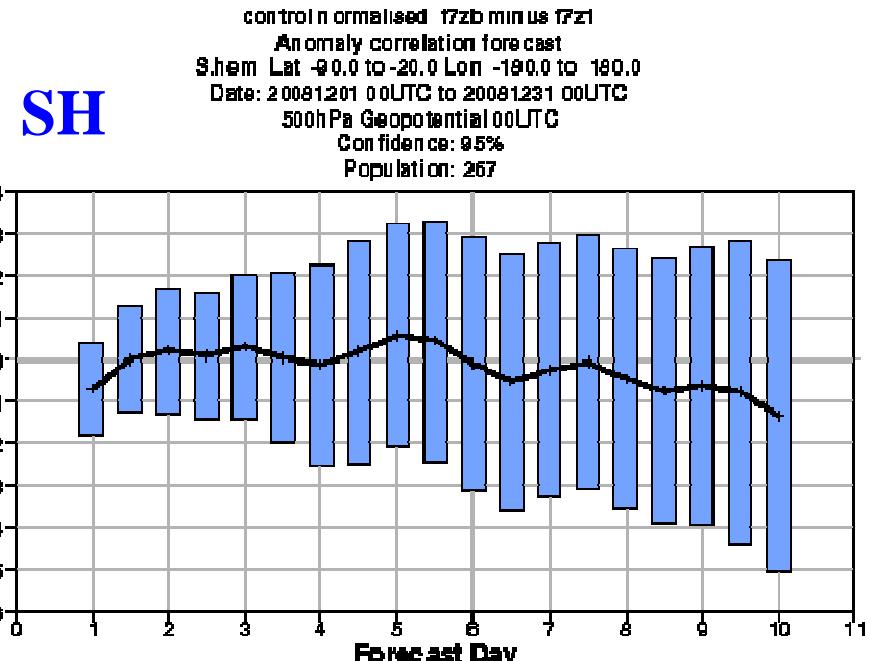
U Tendencies (m/s/day) July from non-oro GWD



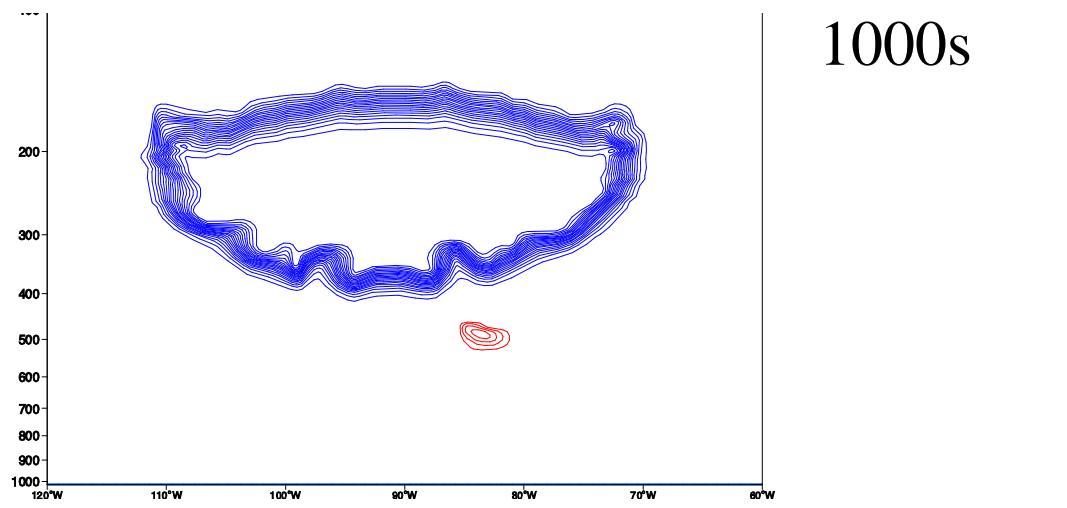
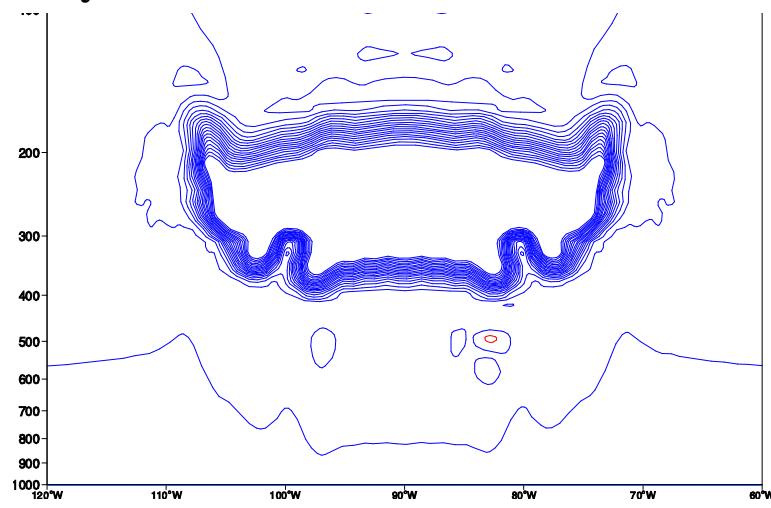
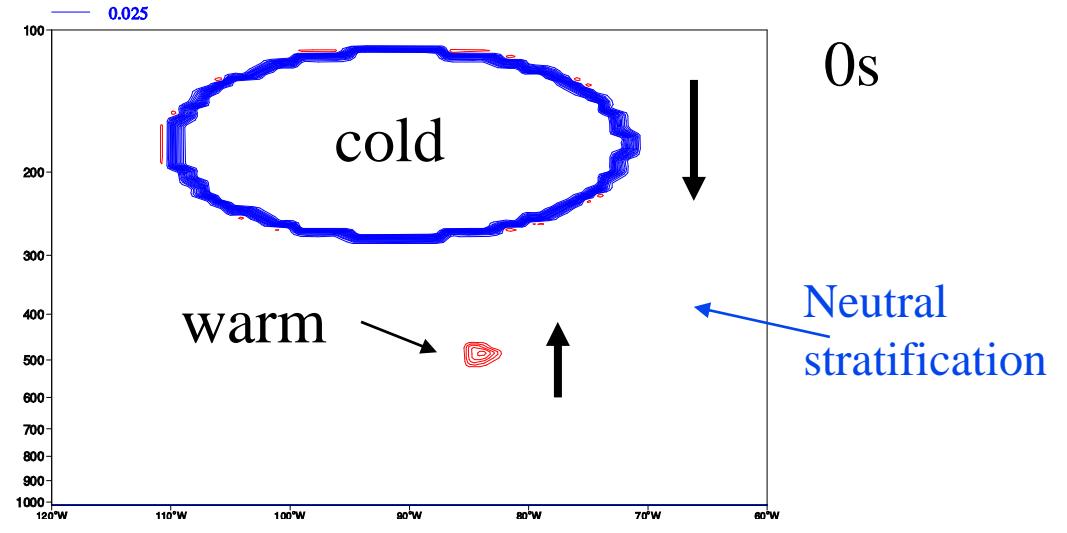
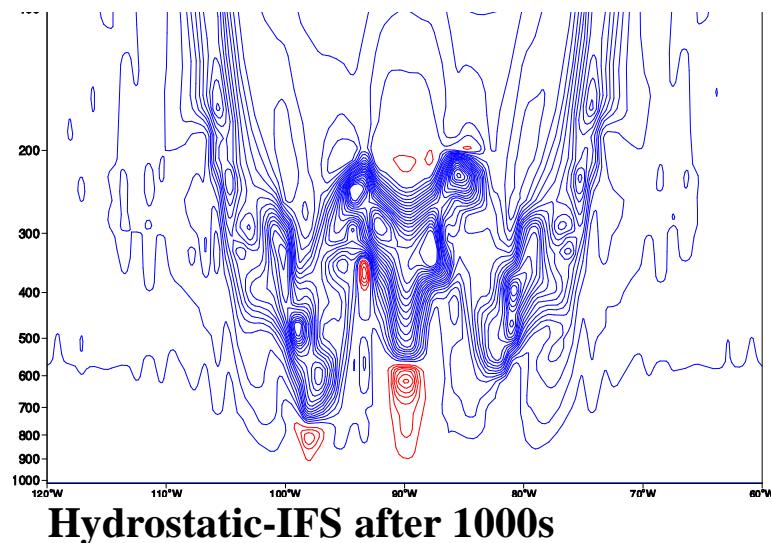
Resolution dependence Cy35r3 August



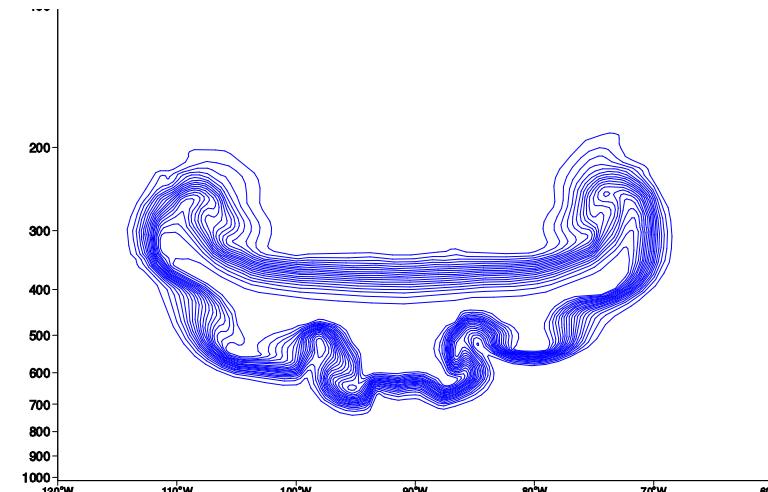
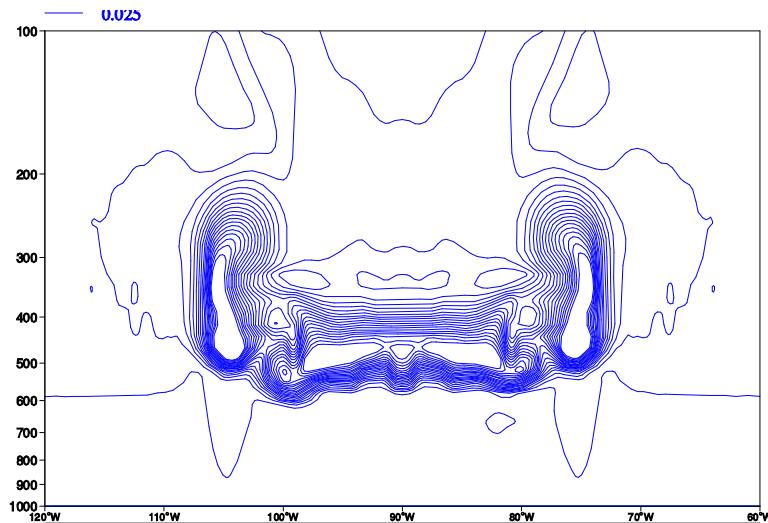
◆ Influence less pronounced over SH as there is no orography, major driver in resolution upgrade !



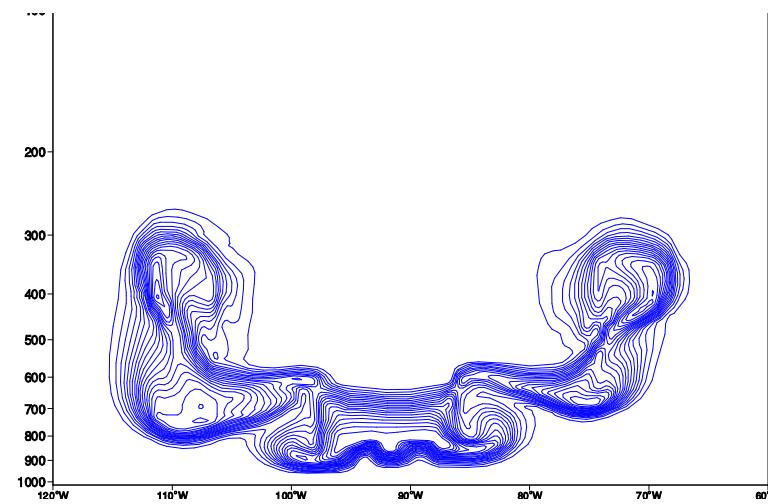
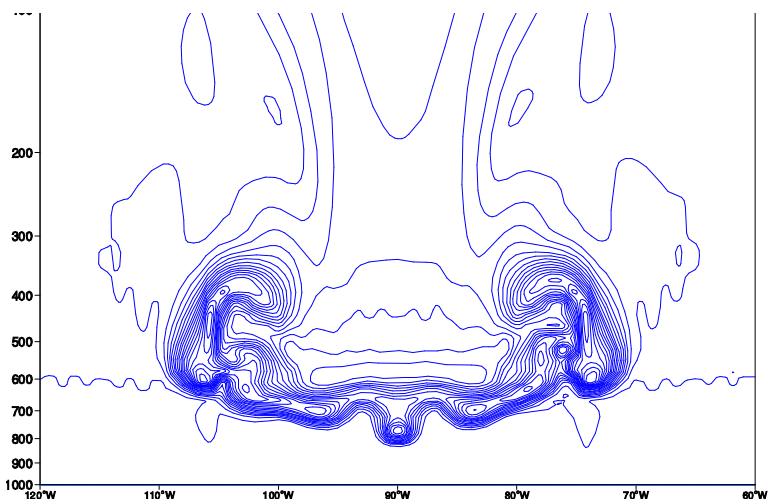
Convective motion (3D bubble test)



Convective motion (3D bubble test)



1800s



2400s

NH-IFS

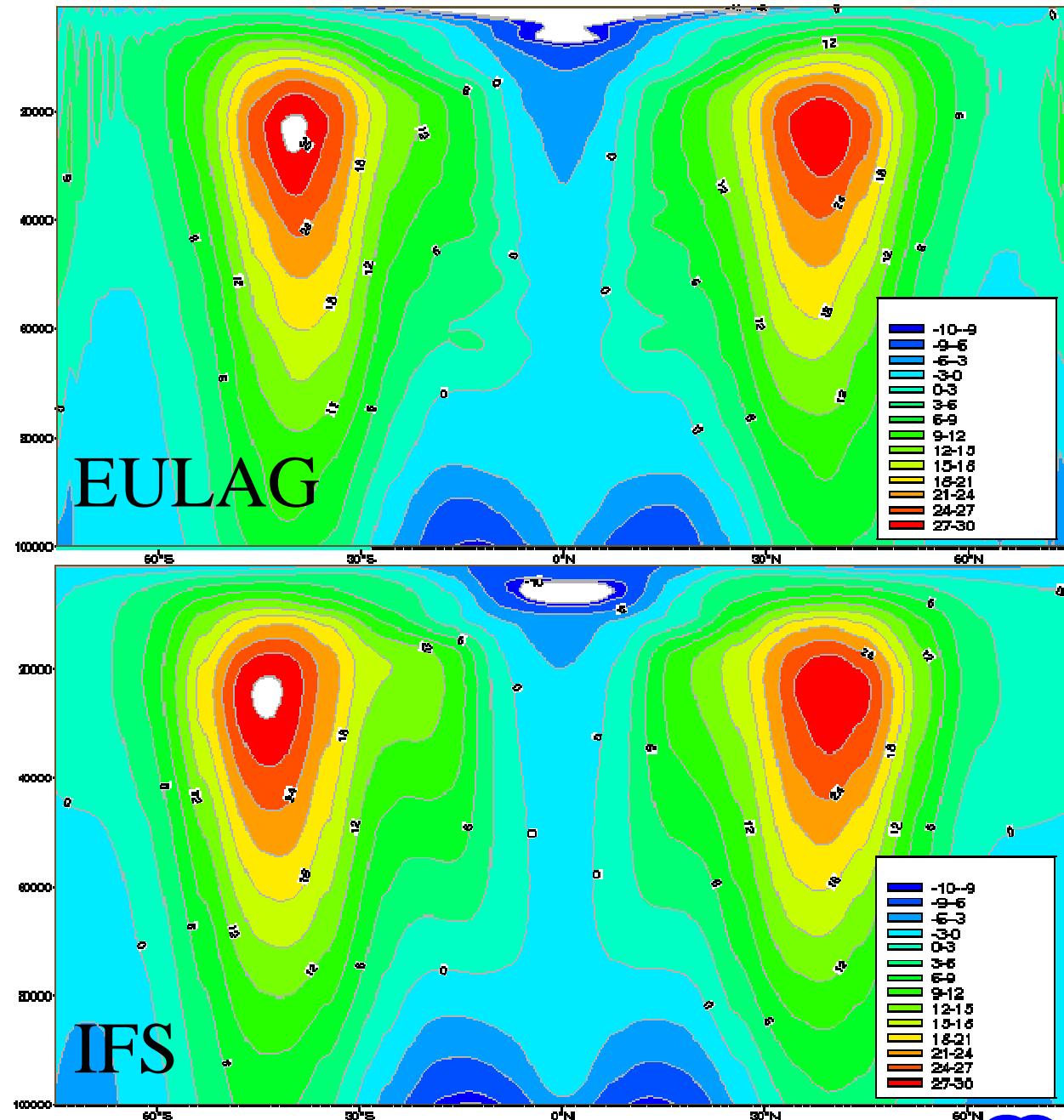
EULAG

Held-Suarez climate

$$a = a_{\text{Earth}} / 10$$

$$\Omega = 10 \times \Omega_{\text{Earth}}$$

zonal-mean
zonal wind U



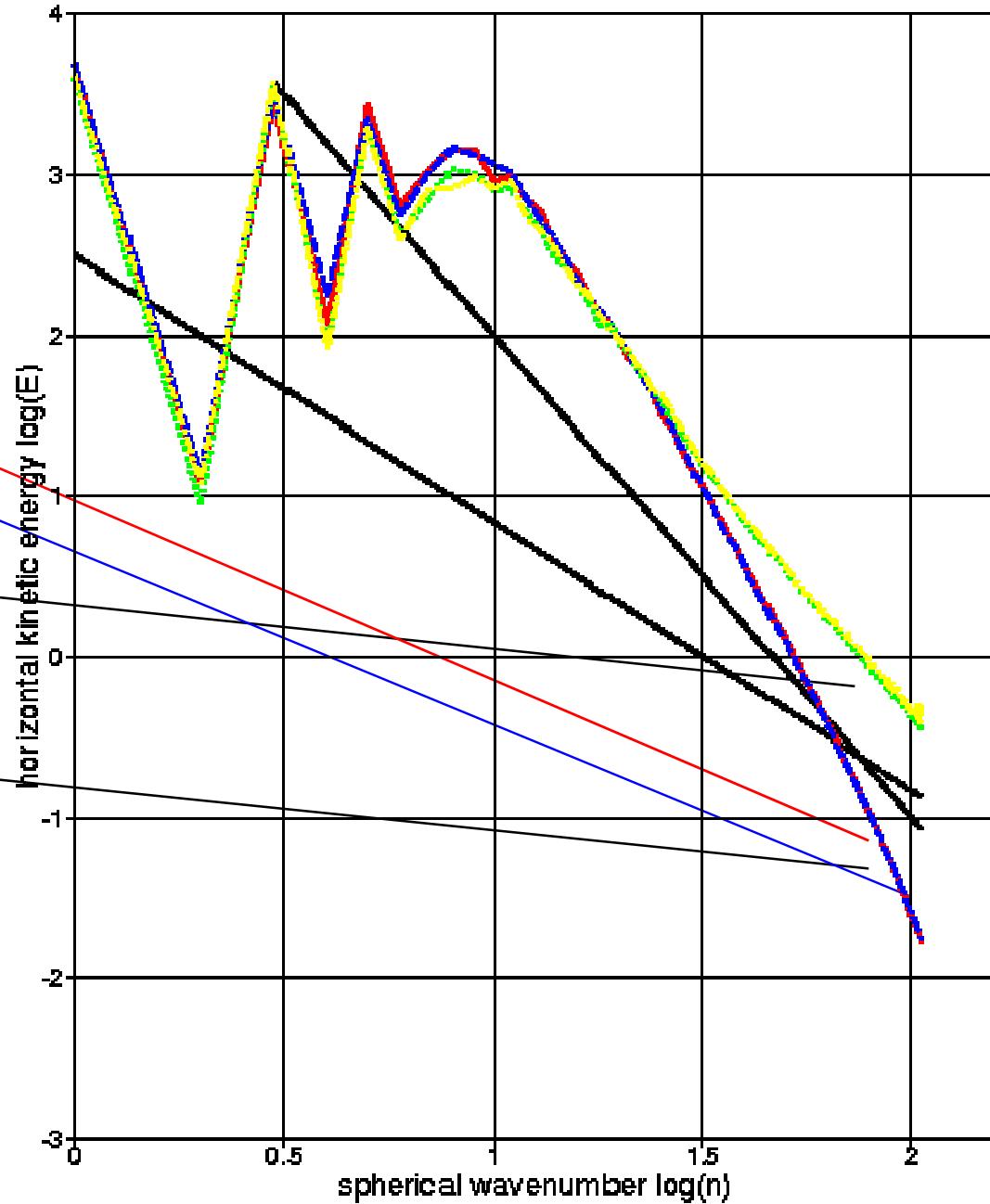
Held-Suarez climate

$$a = a_{\text{Earth}} / 10$$

$$\Omega = 10 \times \Omega_{\text{Earth}}$$

EULAG

IFS

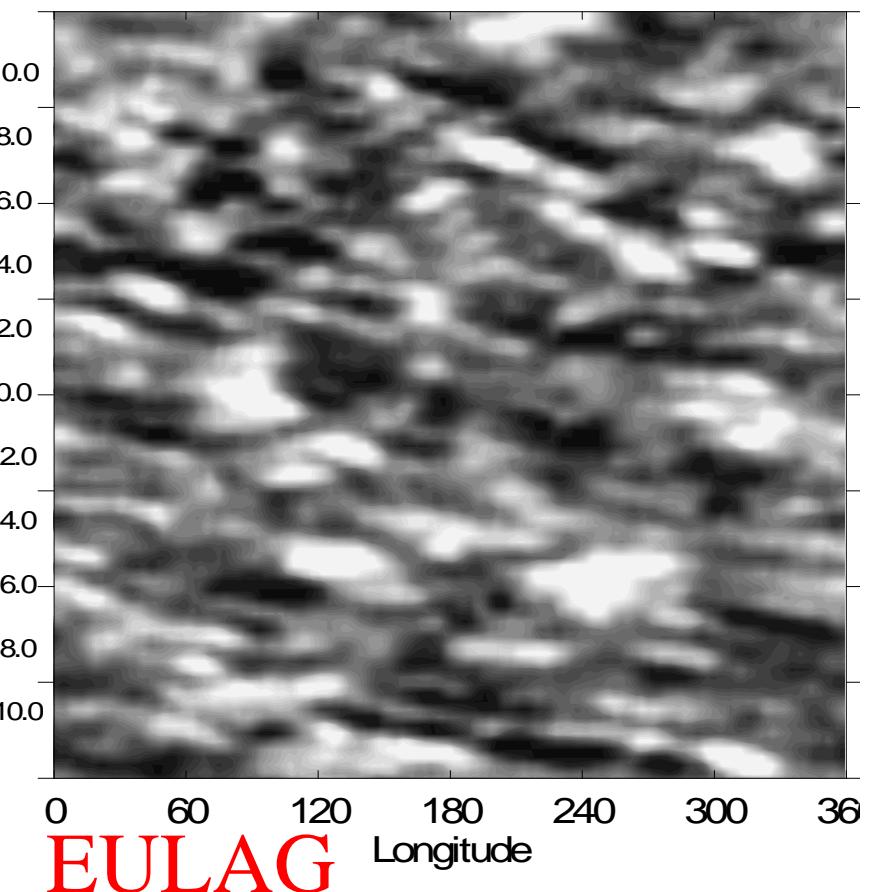
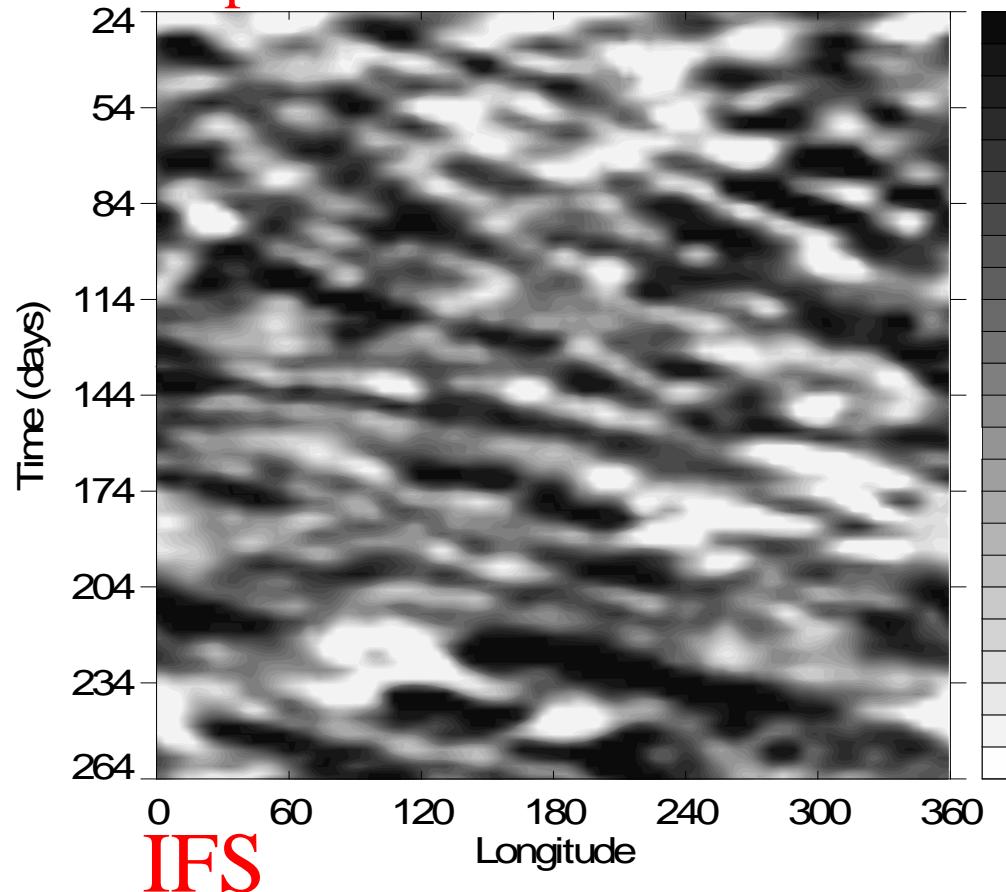


Held-Suarez climate

$a = a_{\text{Earth}} / 10$

Averaged $30^{\circ}\text{N}-50^{\circ}\text{N}$

Temporal anomalies of zonal wind

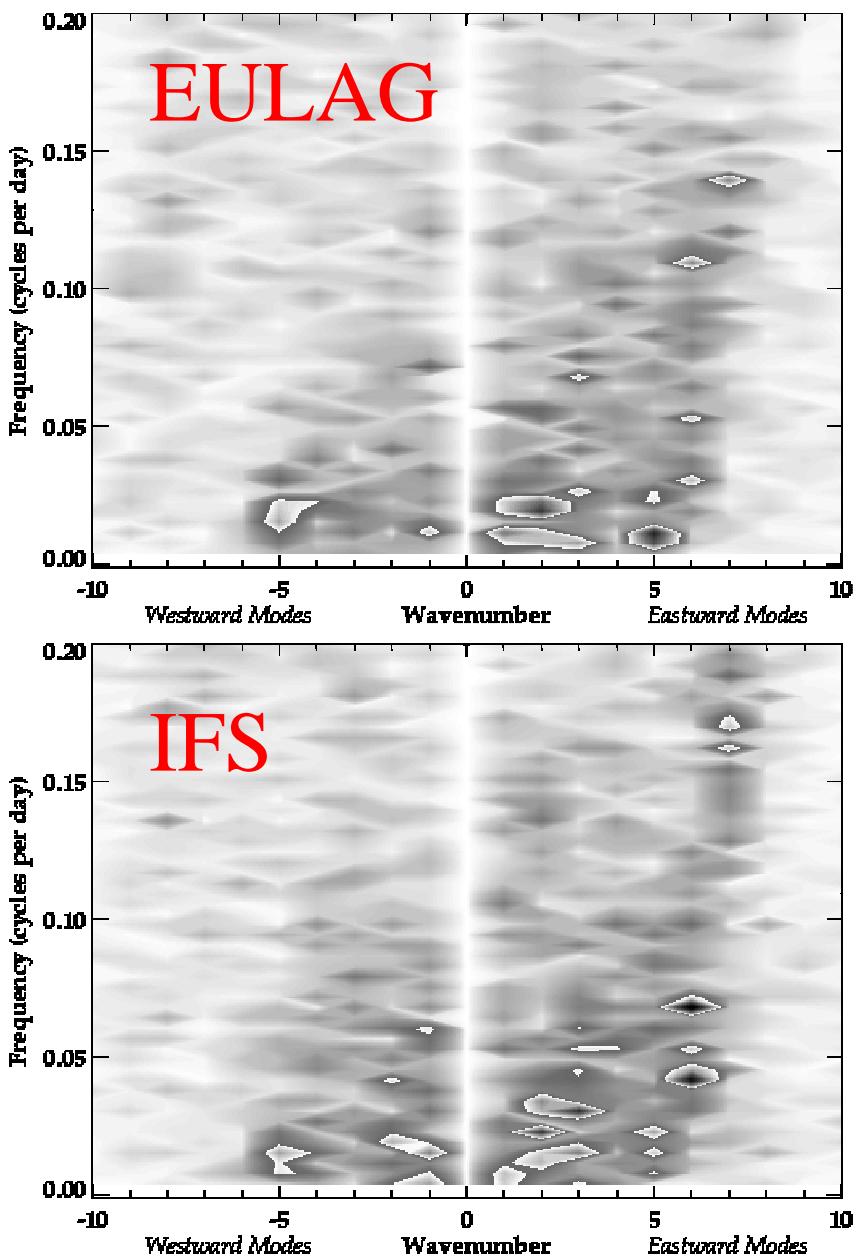


Held-Suarez climate

wavenumber-frequency
diagrams of zonal wind

e.g. Rossby wave dispersion:

$$\omega_R \approx \frac{2\Omega N^2 m}{n(n+1)N^2 + f^2 a^2 (\delta_A N^2 / c_s^2 + k_z^2 + \Gamma_A^2)}$$



Idealized dry atmospheric variability and mean states

- ◆ Zonal mean states compare well with EULAG.
- ◆ **Concern:** Anomalous variability and wavenumber-frequency diagnostics very different !
- ◆ **Concern:** Held-Suarez spectra from wavenumber >20 considerably less energetic in horizontal modes compared to EULAG.

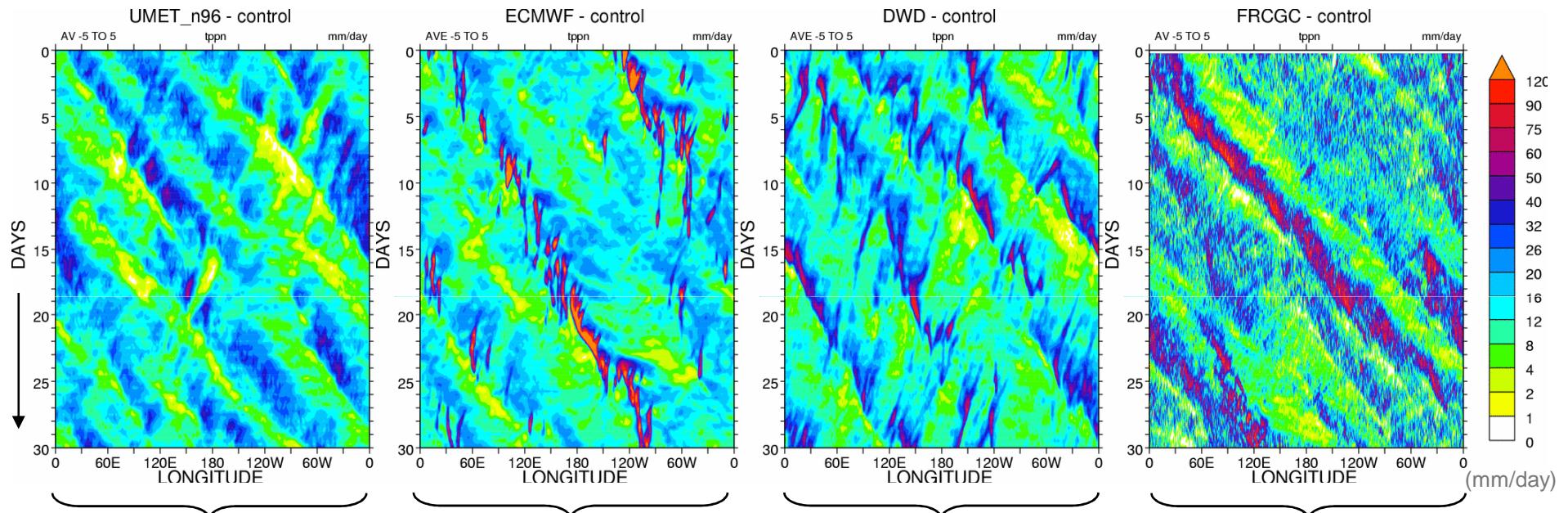
Variability of organized convection

- ◆ The formation of organized convection is sensitive to explicit or implicit viscosity in under-resolved simulations.
- ◆ The sensitivity is consistent with linear theory adapted to include the effect of anisotropy of viscosity (horizontal vs. vertical) at moderate Rayleigh number.

(Piotrowski et al., *J. Comput. Phys.* 2009)

- ◆ It suggests a careful control of the effective numerical viscosity in the numerical core ! (such as the dependence of the truncation error on the derivatives of the flow variables rather than the flow variables themselves).

Higher resolution models average 5°N-5°S



pre-HadGAM1
N96 L38
 $1.25^\circ \times 1.875^\circ$

IFS Cy29r2
 $T_L 159$ L60
 $\sim 125\text{km}$ grid

GME
Icosahedral L31
 $\sim 100\text{km}$ grid

NICAM
Icosahedral L54
7km grid

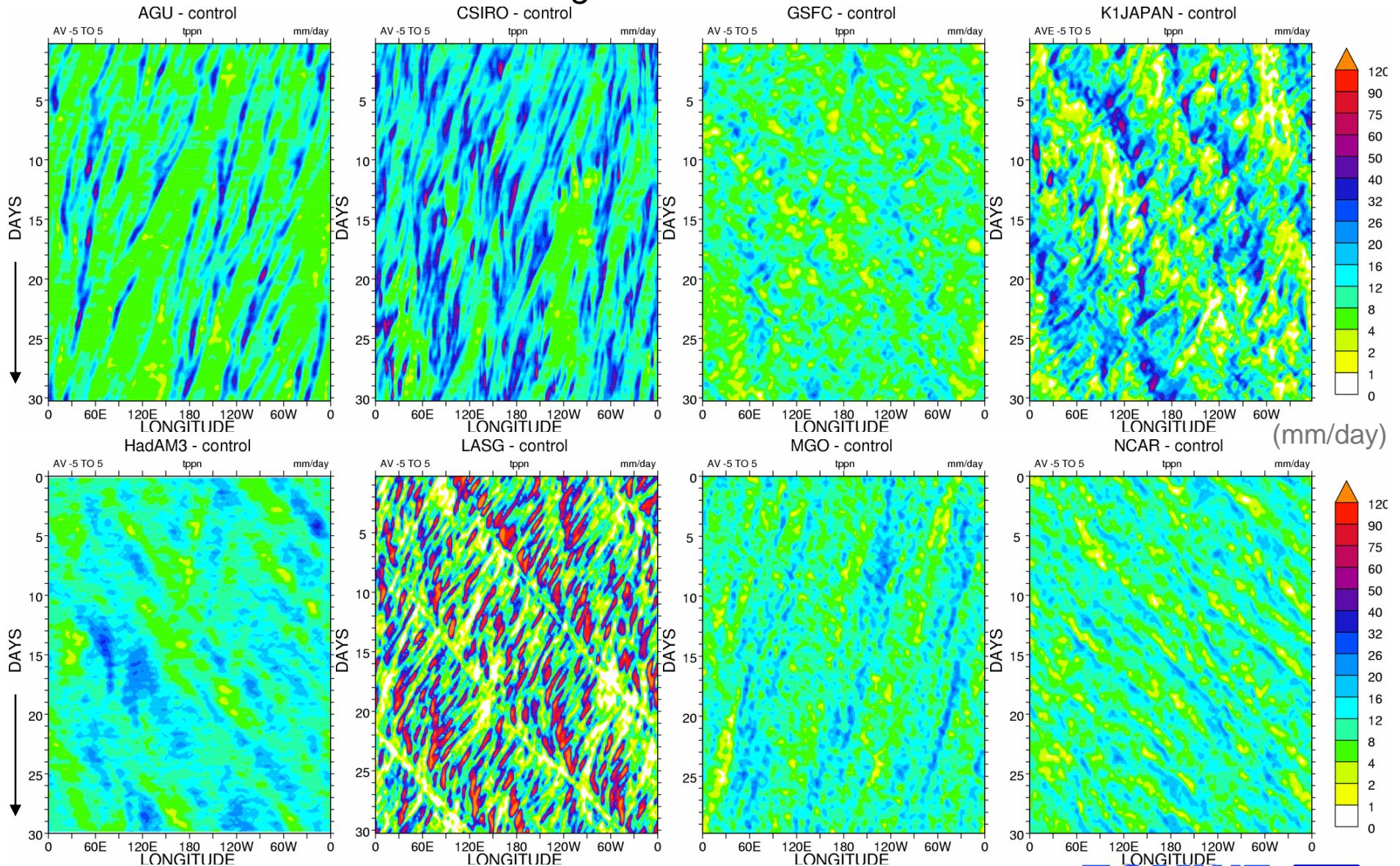
non-hydrostatic
no convective param.

Tropical Variability (precipitation)

Courtesy of D. Williamson



average 5°N-5°S

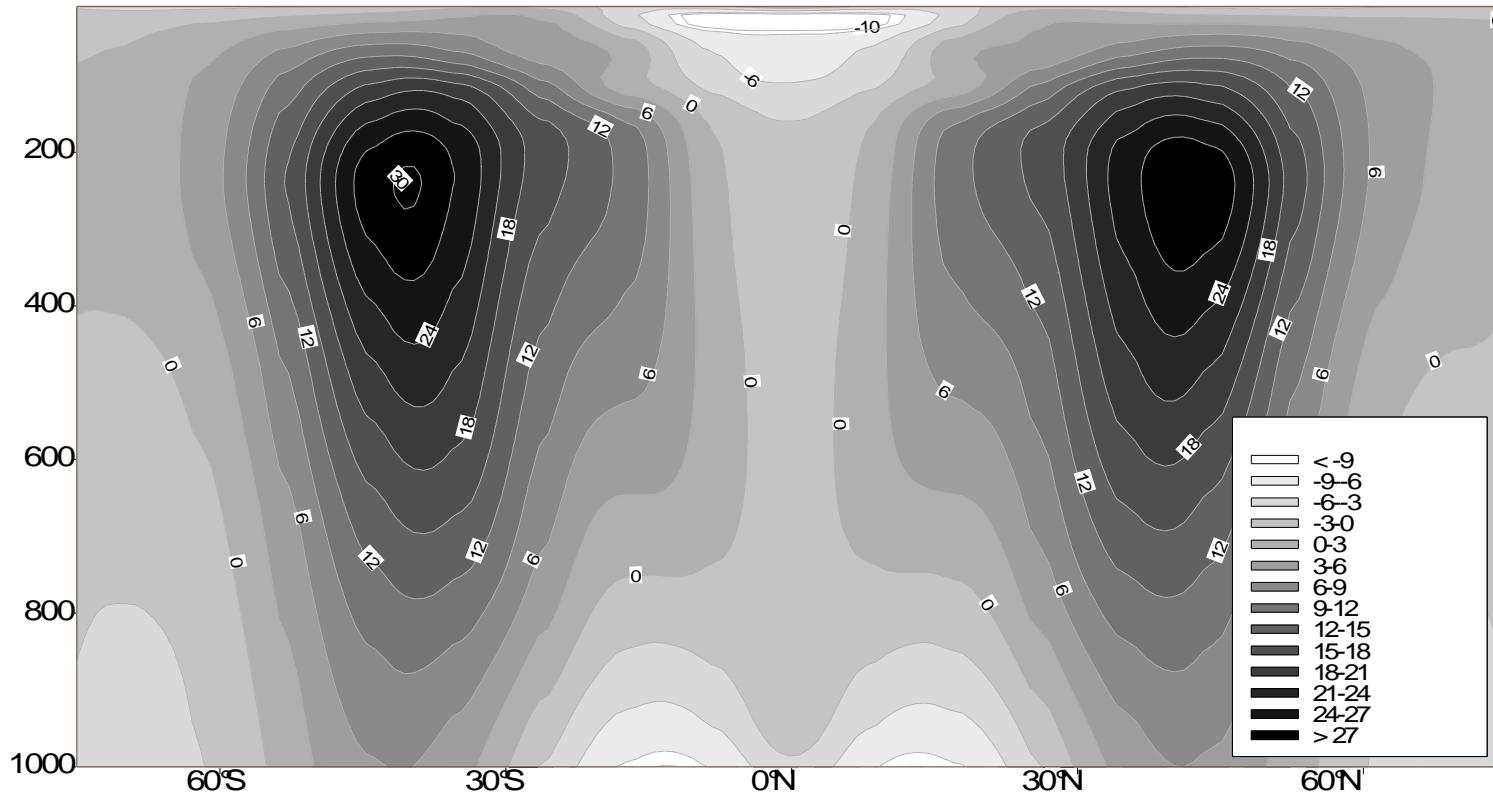


Deep vs. shallow atmosphere

- ◆ Deep atmosphere formulation implemented by Karim Yessad (testing/bugfixing when he visited ECMWF; some remaining issues) (*Staniforth, Wood QJRMS Vol. 129, 1289-1300, 2003*)

The influence of the shallow vs deep atmosphere model formulation

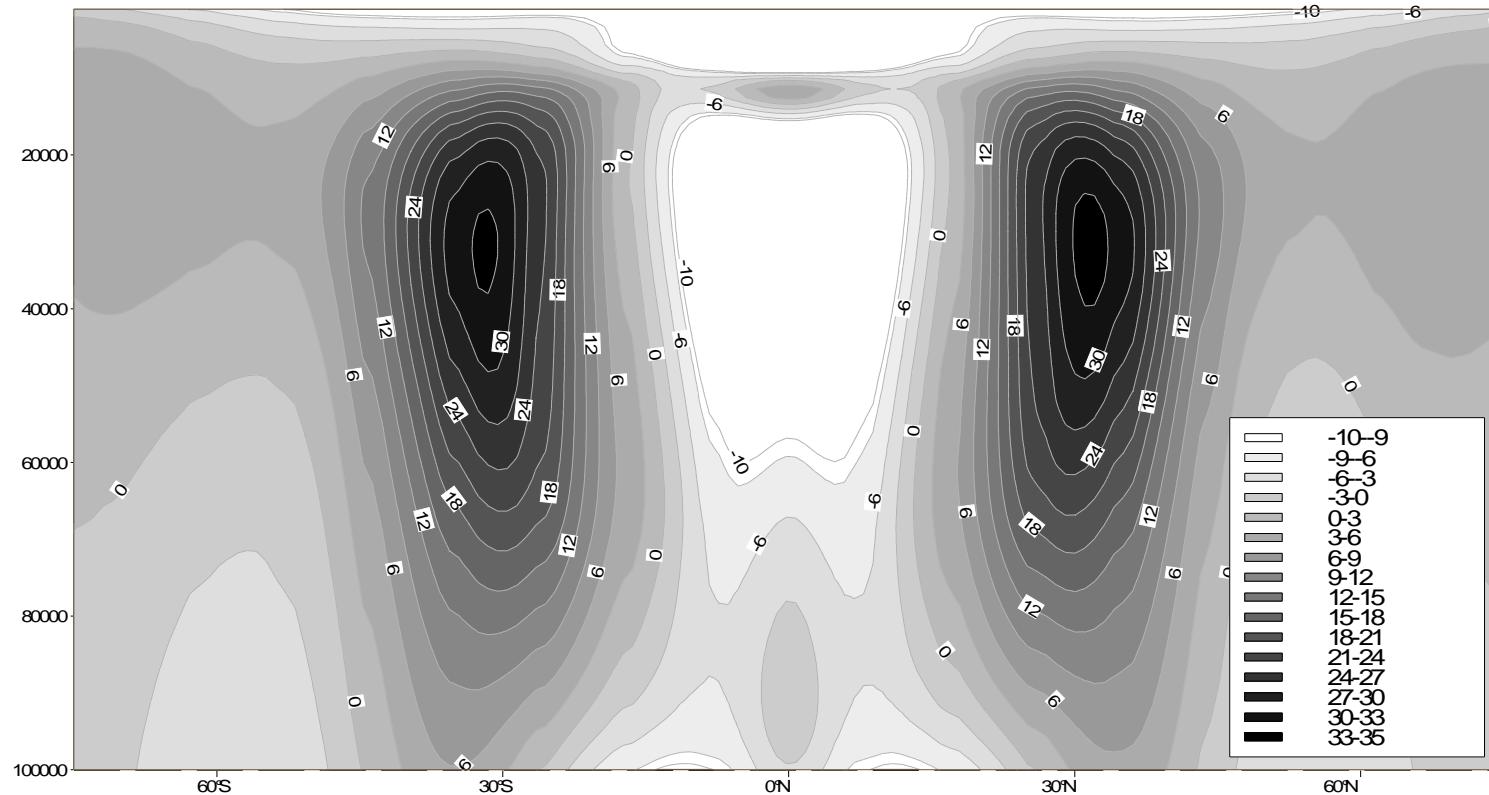
$$\Omega = 20 \times \Omega_{Earth} \quad a = a_{Earth} / 20$$



IFS Held-Suarez simulation with a shallow hydrostatic model
Simmons and Burridge 1981

The influence of the shallow vs deep atmosphere model formulation

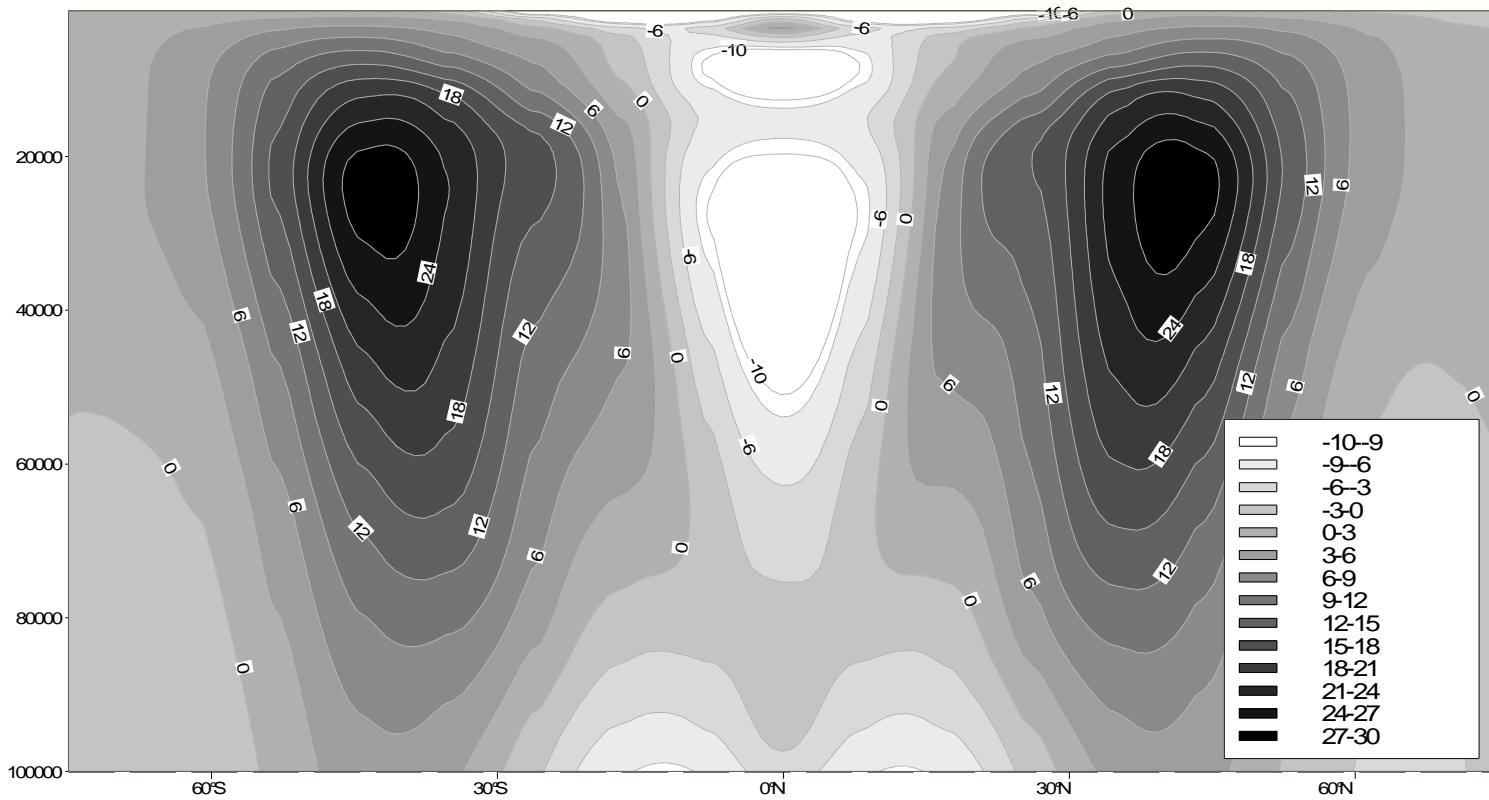
$$\Omega = 20 \times \Omega_{Earth} \quad a = a_{Earth}/20$$



IFS Held-Suarez simulation with a deep hydrostatic model
following *White and Bromley, 1995*

The influence of the shallow vs deep atmosphere model formulation

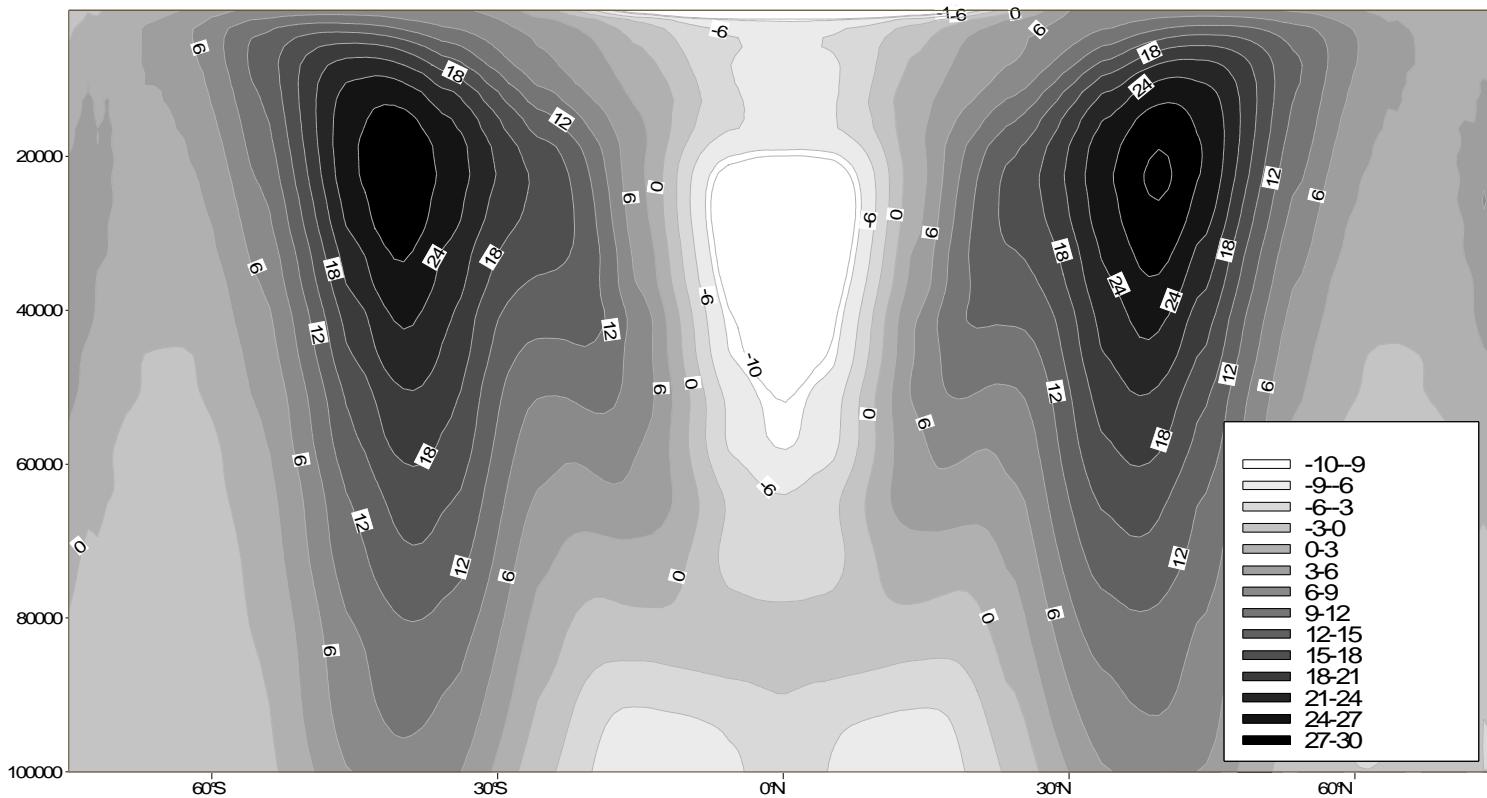
$$\Omega = 20 \times \Omega_{Earth} \quad a = a_{Earth} / 20$$



IFS Held-Suarez simulation with a deep non-hydrostatic elastic model following *Wood and Staniforth 2003*

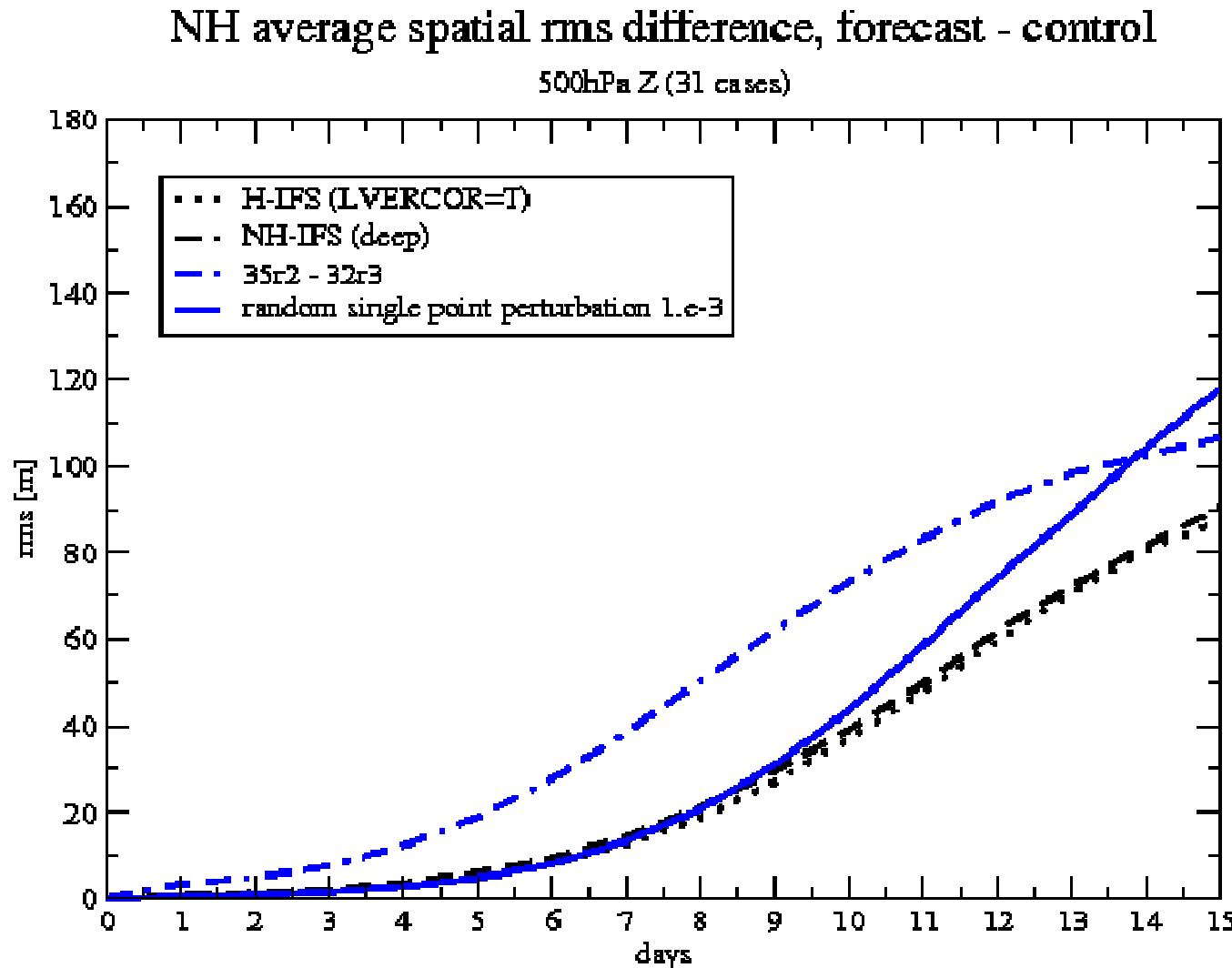
The influence of the shallow vs deep atmosphere model formulation

$$\Omega = 20 \times \Omega_{Earth} \quad a = a_{Earth} / 20$$



EULAG Held-Suarez simulation with a deep non-hydrostatic
anelastic model, Wedi and Smolarkiewicz, 2009

The influence of the H-shallow vs NH-deep atmosphere model formulation



Seasonal climate

- ◆ ***Shallow hydrostatic and nonhydrostatic*** seasonal climate simulations are **the same** for T_L159L91 (13 month) 4 member ensemble except in the stratosphere (LVERTFE=F)
- ◆ ***Deep hydrostatic and deep nonhydrostatic*** seasonal climate simulations are **very similar** to the ***shallow hydrostatic*** simulations. Some warming of the polar regions and a cooling of the upper stratosphere can be noticed compared to the control.