

Republic Hydrometeorological Service of Serbia (establishted 1888, WMO member since 1947)



## Modelling cloud-aerosol (dust) interactions: a potential for further NWP improvements

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# Why to consider atmospheric desert dust as a factor in NWP models?

- Atmosphere is modified by dust
  - Through <u>direct effects</u> (affecting radiation)
  - Through <u>indirect effects</u> (affecting clouds)
- Why dust is a major aerosol?
  - Most abundant
  - Specific chemical and physical features favorable for direct and indirect effects

**IPCC:** Both magnitude and the sign of dust radiative forcing yet unresolved (unknown positive or negative)



## How much dust affects the atmospheric radiation?

The outgoing longwave radiation anomaly in the UK operational model over the Sahara due to dust

(Haywood, et al., 2003)







### Cooling surface temperature by ~5°C in DREAM model (not only over Sahara!) (Nickovic et al, 2004; Perez et al, 2006)

## Heterogeneous cold clouds formation

- Several *Science* and *Nature* articles published since 2013 indicate the importance of dust speciffically
- Mineral dust particles act as <u>the most efficient</u> heterogeneous ice nuclei in the tropospheric clouds
- Dust particles lifted to the colder tropopause cause earlier glaciation of supercooled cloud water



### Ice formation and precipitation

#### Koop and Mahowald, Nature, 2013

## Dust: key catalyst for cold-cloud formation even far away from sources

- 2/3 of ice clouds formed due to pure dust and dust metalics
- Only small dust concentration needed
- Dust mineralogy matters!

#### Sciencexpress

#### Clarifying the Dominant Sources and Mechanisms of Cirrus Cloud Formation

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## **DREAM - Dust Regional Atmospheric Model**

(Nickovic et al, 2001; Pejanovic et al, 2010; Vukovic et al, 2013)



- Widely used dust model in the community
- Operational dust forecasts within the WMO dust SDS-WAS model intercomparison project



### **Dust data assimilation in DREAM**

(Nickovic, Pejanovic, Solonos, Cvetkovic, Petkovic, work in progress

- Collaboration with NOA (Greece) and UK MetOffice
- Observations: MSG/SEVIRI Dust Optical
   Depth (DOD) over ground only for the
   moment

 $+k(C-C_T)=0$ 

- Newtonean Nudging  $k \sim 10^{-5}$ 

 $k \sim 10^{-4}$ 

k≈10<sup>-3</sup>

6

С





## **IN parameterization in NMM-DREAM**

Example of a typical cloud parameterization in today's models

$$\frac{\# IN = 100 \ m^{-3} = const}{\partial t} \quad \frac{\partial N_{ice}}{\partial t} = \dots - \frac{N_c (\# IN)}{q_{ice}} ICEGEN \quad \frac{\partial q_{ice}}{\partial t} = \dots - f (ICEGEN)$$

e.g. Bangert et al, 2011

- Most operational microphysics schemes use predefined #IN
- Instad, we plan to use #IN as predicted variable in the Thompson "dust-friendly" microphysics (MWR,2012)



## **DREAM #IN parameterization**

#### **DREAM dust model**

- 25km resolution; Sahara/Mediterranean region
- $-\,$  Particle bin radii: 0.15, 0.25, 0.45, 0.78, 1.3, 2.2 ,3.8 , 7.8  $\mu m$

Immersion ice nucleation (two options) [-35°C <T<-5°C]

DeMott et al, (2010)

$$n_{IN} = (n_{dust})^a 10^{[bT+c]} \left| \frac{\#}{m^3} \right|$$

Niemand et al (2012)

$$n_{IN} = S_{dust} e^{-mT+n} \left[\frac{\#}{m^3}\right]; S_{dust} = surface of dust particles$$

**Deposition ice nucleation [-60°C <T<-35°C]** 

Steinke al (2014)

$$n_{IN} = S_{dust} 1.88 \cdot 10^5 e^{-pT + q(RH_{ice} - 100\%)} \left[\frac{\#}{m^3}\right];$$

## Model #IN vs. MPL lidar, Izana





0.0

MPL-3 S.C. de Tenerife 2013-08-21



MPL-3 S.C. de Tenerife 2013-08-22



20

MPL Lidar (Tenerife)

1.5

0.0

2.0

1.5

0.5



## September 2012 event

• One week of moderate Saharan dust in the central Mediterranean











MODIS maps produced with the Giovanni, developed and maintained by the NASA GES DISC

### Sep 2012 dust case - Potenza

z-t graph
Model #IN (color bar)
vs.
MIRA55 Ice Cloud

Water(black line)



## Thank you !